



- **Announcements**

- **HW 2**

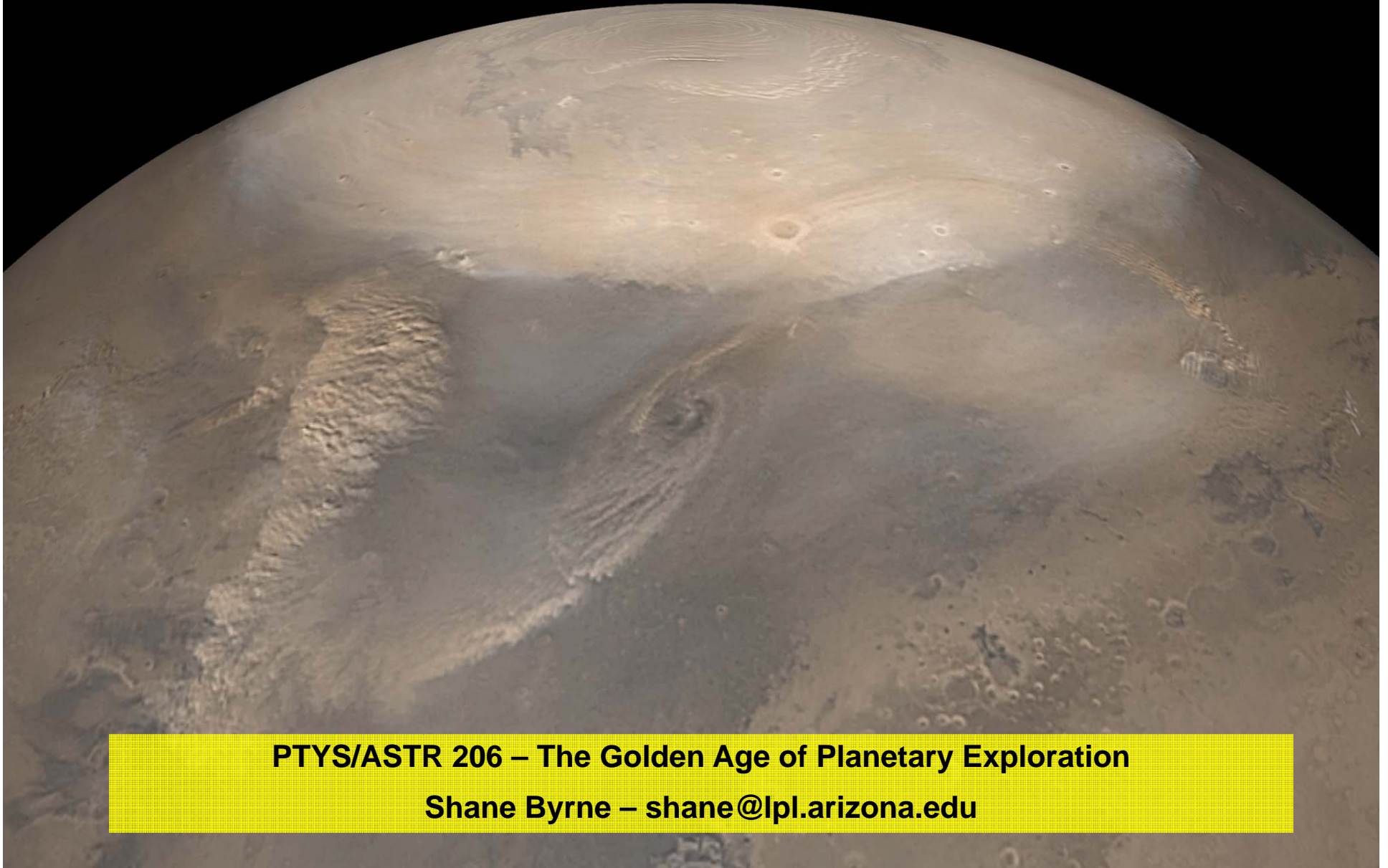
- **Mid-term results**

- **Mid-term went very well**

- **Results**

- **87.5 – 100%**                      **17**
- **70-87.5%**                         **37**
- **62.5-70%**                         **29**
- **50-62.5%**                         **19**
- **<50%**                                **21**

# Terrestrial planet atmospheres



PTYS/ASTR 206 – The Golden Age of Planetary Exploration

Shane Byrne – [shane@lpl.arizona.edu](mailto:shane@lpl.arizona.edu)

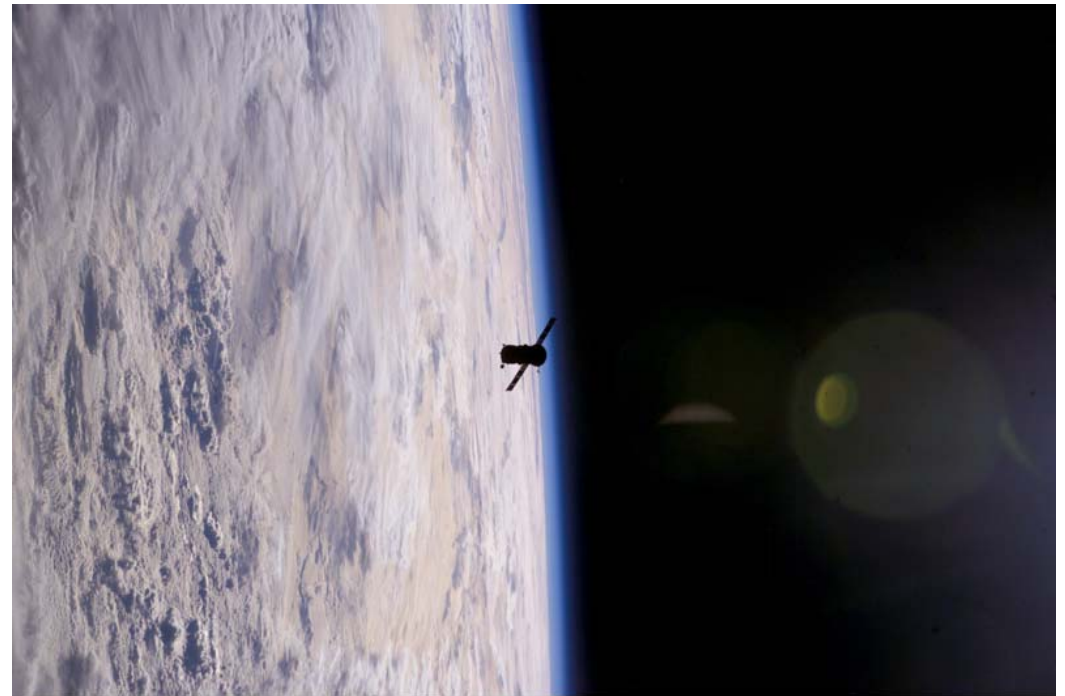
## In this lecture...

- **Introduction to atmospheres**
  - Pressure and Temperature
  - Scale height
  - Comparing planetary atmospheres
- **Radiation and atmospheres**
  - Clouds
  - Greenhouse effect
- **Circulation**
  - Why is Tucson a Desert?
  - What's El Nino
- **Surface features from the atmosphere**
  - Sand dunes and how they work
  - Eolian erosion
- **'Atmospheres' of Moon and Mercury**
  - Ice in polar craters

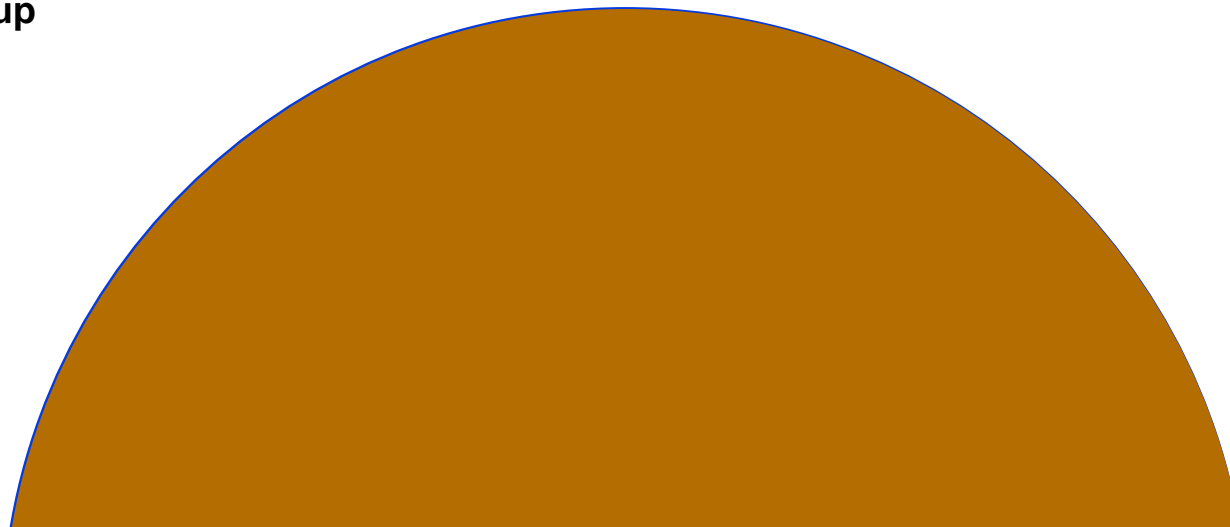


## Introduction

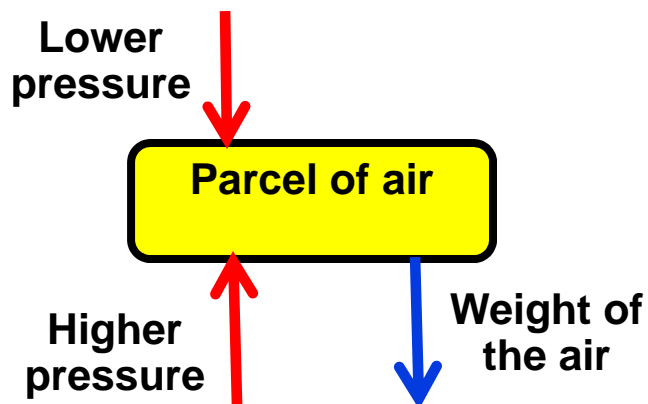
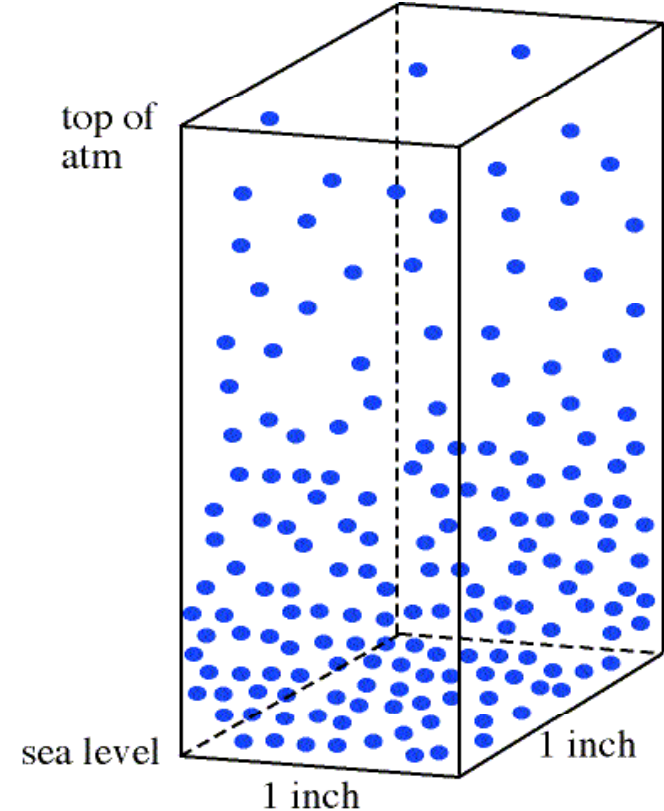
- What's the atmosphere?
  - A thin gas layer
  - Gas molecules moving around at high-velocities
  - Held down with gravity
- The “edge” of the atmosphere
  - Thins gradually with altitude
  - Karman line
    - ◆ Based on aeronautics
    - ◆ ~100km up



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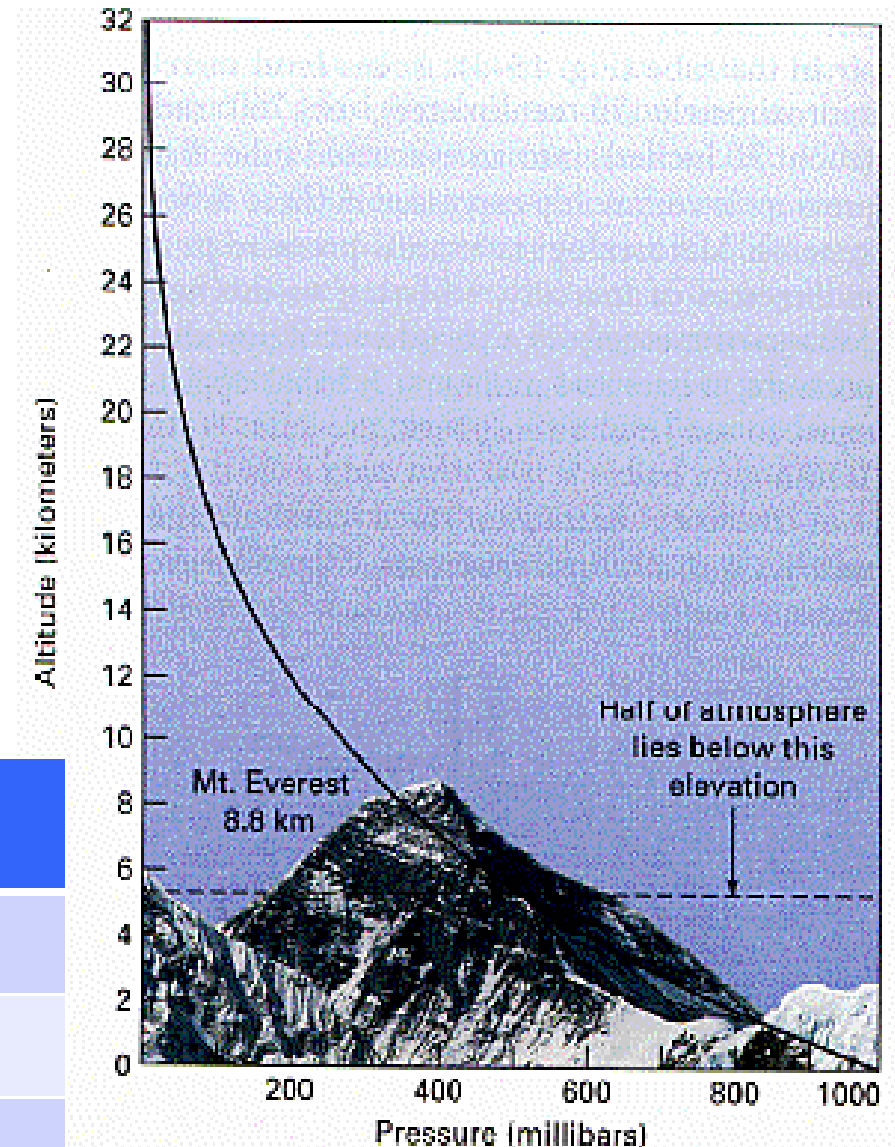
- **Gravity squeezes the atmosphere**
  - High pressure at the bottom
  - $10^5$  Pascals (Newtons per square meter)
  - ~10,000 Kg of mass above each square meter
  - So why aren't we squashed flat?
- **The atmosphere is supported by pressure**
  - Pressure =  $a \cdot \text{density} \cdot \text{temperature}$ 
    - ◆ 'a' depends on the type of gas
  - A cold dense atmosphere (e.g. Titan) can have the same pressure as a hot less-dense atmosphere (e.g. Earth)



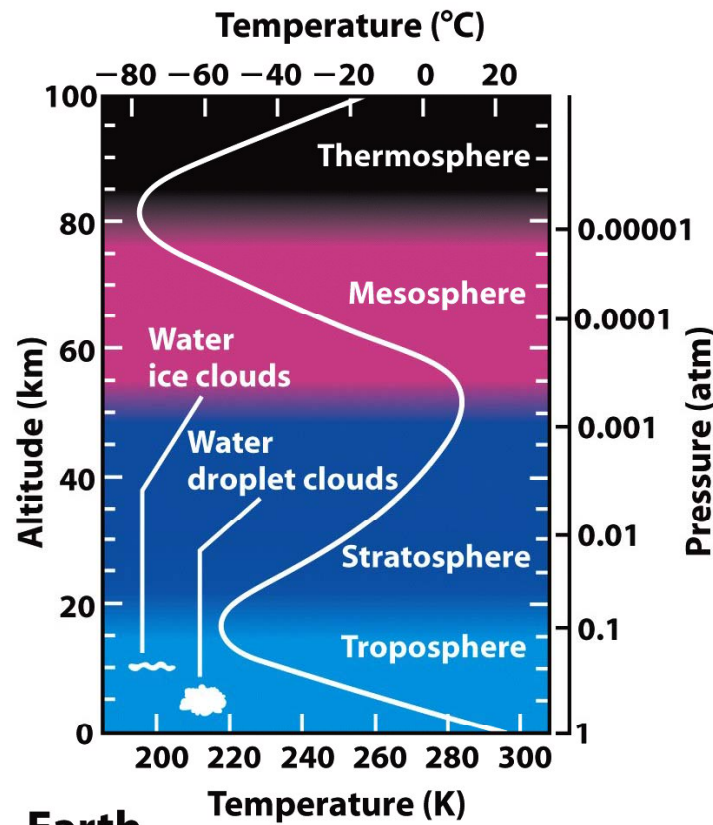
**This is “hydrostatic equilibrium”  
Makes the atmosphere (mostly) stable**

- Pressure drops with height
- The ‘scale height’ measures the thickness of the atmosphere
  - At the scale height the pressure is  $1/e$  times the surface pressure.
  - $e$  is a special number in math
  - $1/e = 37\%$
  - Earth’s scale height is  $\sim 8\text{km}$
- Scale height tells you how compact the atmosphere is

Scale height	Elevation	Atm. Pressure
1	8km	$37\% = (1/e)^1$
2	16km	$14\% = (1/e)^2$
3	24km	$5\% = (1/e)^3$

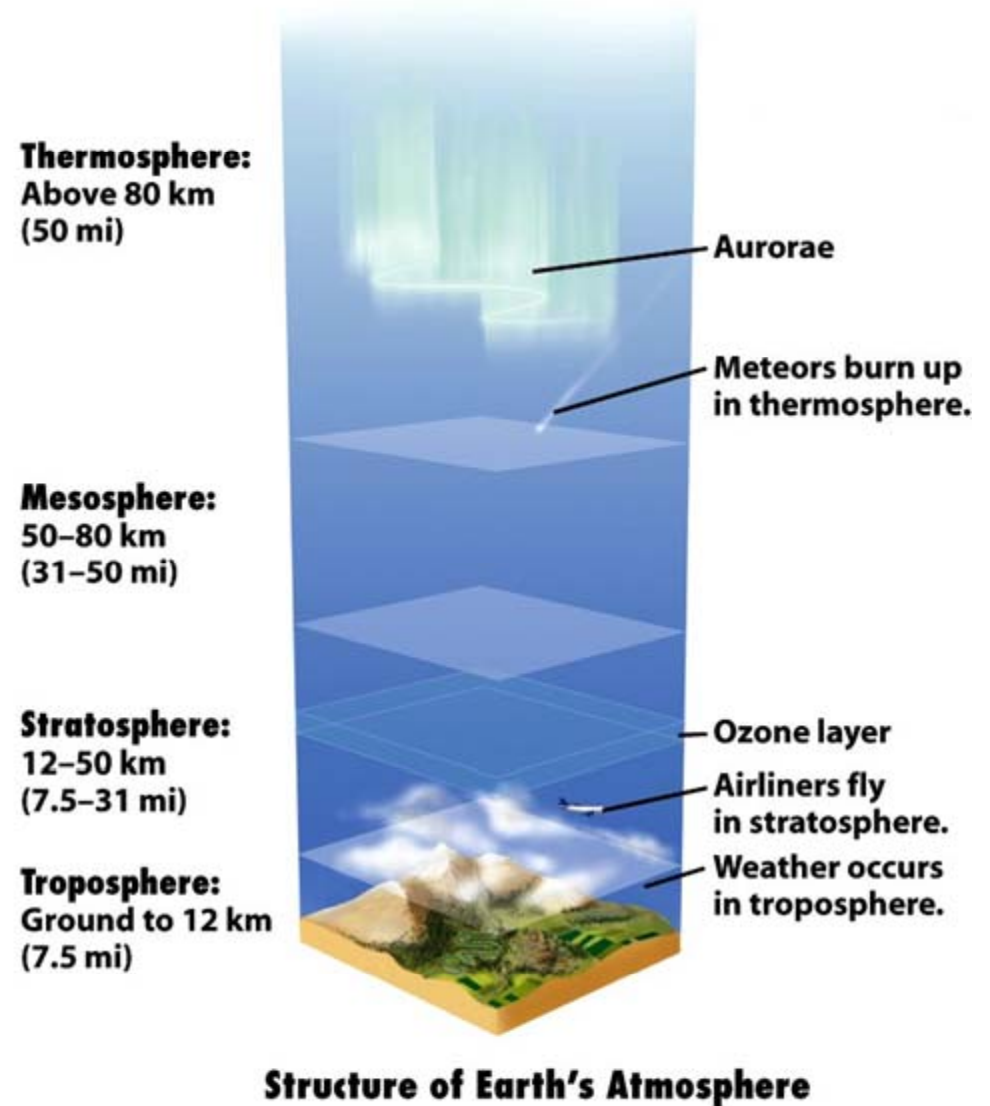


- Earth's atmosphere has many sections
- All our surface features (even the highest mountains) are in the troposphere



Earth

Figure 11-27a  
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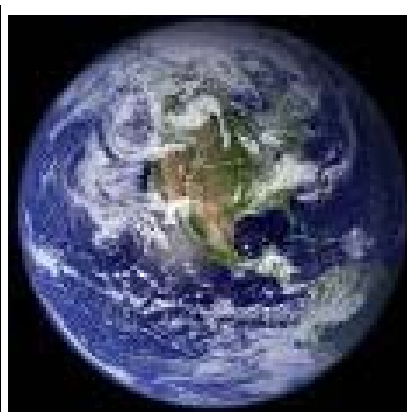
Structure of Earth's Atmosphere



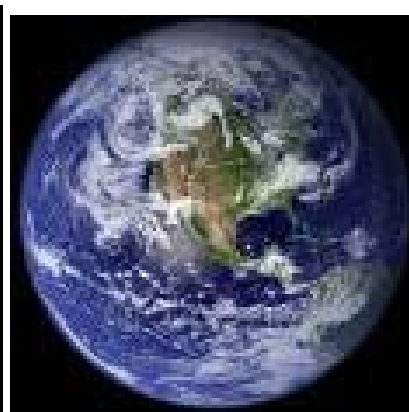
- **Here's a problem**
  - **What's the atmospheric pressure at the Karman line – 100km elevation**
  - **100km is 12.5 scale heights (i.e. 100km / 8km)**
  
  - **The pressure is  $(1/e)^{12.5}$**
  - **Remember  $1/e$  is 0.37 (or 37%).**
  
  - **So  $(0.37)^{12.5}$  is 0.000004**
  - **The atmospheric pressure at 100km (the 'edge' of space) is 4 millionths that at sea-level**



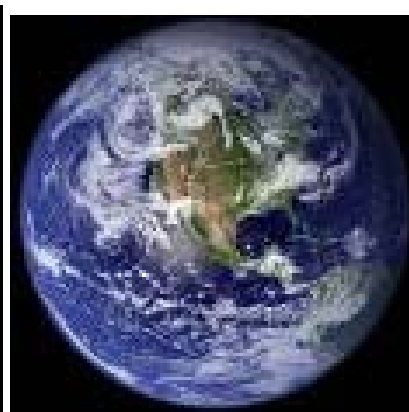
- Different planets have different gases, temperatures and gravities
  - If gravity goes up?
  - If temperature goes up?
  - What if the atmospheric gases were more massive? E.g.  $\text{CO}_2$  vs.  $\text{N}_2$



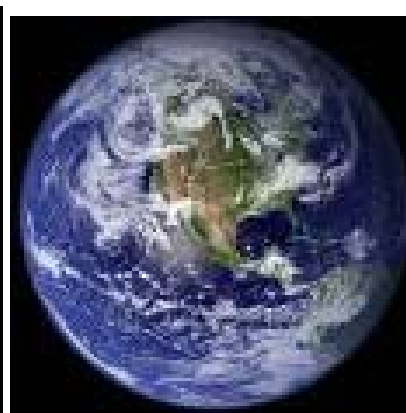
- Different planets have different gases, temperatures and gravities
  - If gravity goes up?
    - ◆ Higher gravity makes the atmosphere more compact
    - ◆ Smaller scale height
  - If temperature goes up?
  - What if the atmospheric gases were more massive? E.g.  $\text{CO}_2$  vs.  $\text{N}_2$



- **Different planets have different gases, temperatures and gravities**
  - **If gravity goes up?**
    - ◆ Higher gravity makes the atmosphere more compact
    - ◆ Smaller scale height
  - **If temperature goes up?**
    - ◆ Higher temperatures makes the atmosphere less compact
    - ◆ Larger scale height
  - **What if the atmospheric gases were more massive? E.g. CO<sub>2</sub> vs. N<sub>2</sub>**



- **Different planets have different gases, temperatures and gravities**
  - **If gravity goes up?**
    - ◆ Higher gravity makes the atmosphere more compact
    - ◆ Smaller scale height
  
  - **If temperature goes up?**
    - ◆ Higher temperatures makes the atmosphere less compact
    - ◆ Larger scale height
  
  - **What if the atmospheric gases were more massive? E.g. CO<sub>2</sub> vs. N<sub>2</sub>**
    - ◆ Heavier gases make the atmosphere more compact
    - ◆ Smaller scale height



- Comparing planetary atmospheres
- Compositions are very different

**Table 9-4** Chemical Compositions of Three Planetary Atmospheres

	Venus	Earth	Mars
Nitrogen (N <sub>2</sub> )	3.5%	78.08%	2.7%
Oxygen (O <sub>2</sub> )	almost zero	20.95%	almost zero
Carbon dioxide (CO <sub>2</sub> )	96.5%	0.035%	95.3%
Water vapor (H <sub>2</sub> O)	0.003%	about 1%	0.03%
Other gases	almost zero	almost zero	2%

### Titan

95%

Zero

Zero

Zero

5% methane



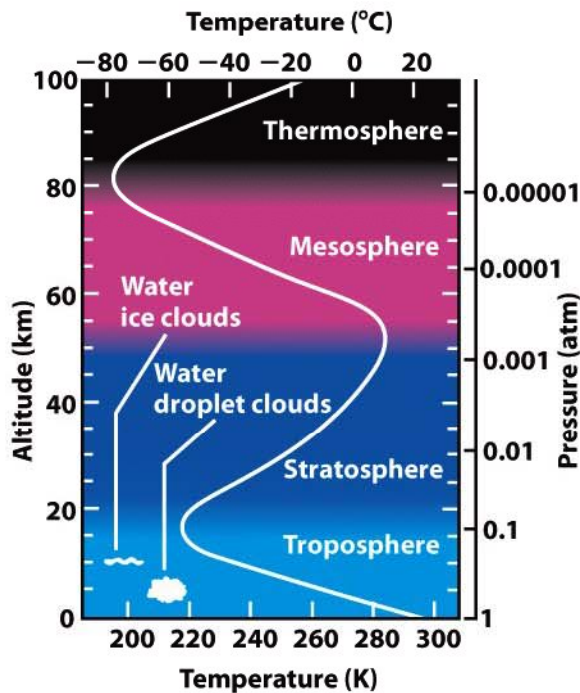


- **Carbon dioxide very common for the inner planets**
  - **Not on Titan or other outer solar system bodies**
  - **Any guesses why?**
  
  - **Not on Earth...**
  - **Any guesses why?**

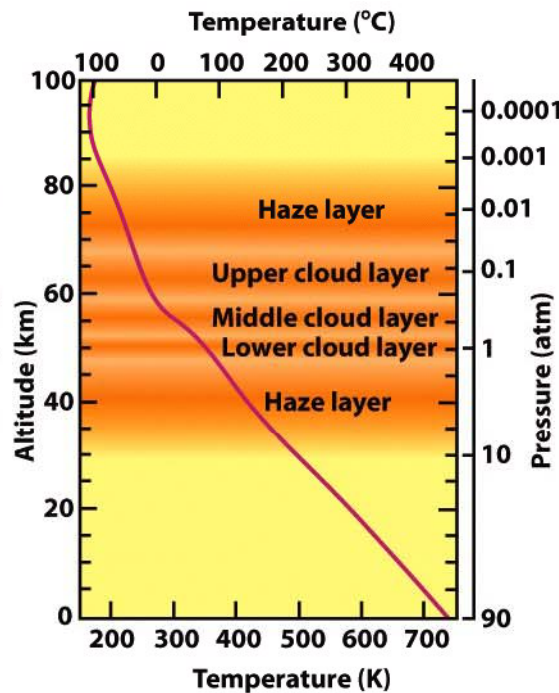


- **Carbon dioxide very common for the inner planets**
  - **Not on Titan or other outer solar system bodies**
  - **Any guesses why?**
  - **Too cold! Carbon dioxide is frozen solid in the outer solar system**
  
  - **Not on Earth...**
  - **Any guesses why?**
  - **Carbon dioxide used to dominate Earth's atmosphere**
  - **Removed by plants**

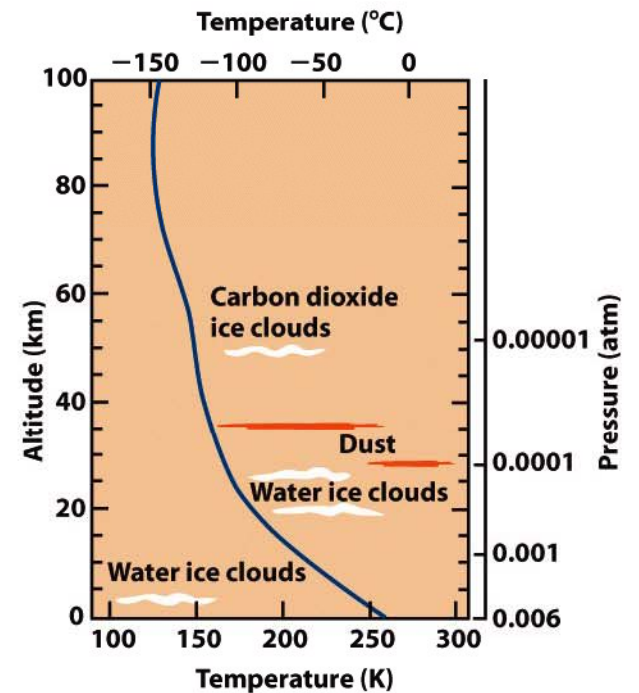
- Comparing the amounts of atmosphere...
- Venus, Earth and Mars are hugely different
  - Earth's surface has a pressure of ( $10^5$  Pa) or 1 bar or 1 atmosphere
  - Venus surface is about 90 bars
  - Mars surface is about 0.006 bars



(a) Earth



(b) Venus



(c) Mars



## Radiation and atmospheres

- **Three major effects**
  - **Reflection from clouds**
    - ◆ Water on Earth
    - ◆ Sulfuric acid on Venus
    - ◆ Cools the surface
  - **Greenhouse effect**
    - ◆ Warms the surface
    - ◆ CO<sub>2</sub> on Earth, Mars and Venus
  - **Transports heat to cold areas**
    - ◆ Polar night on Earth and Mars
    - ◆ Night on Venus

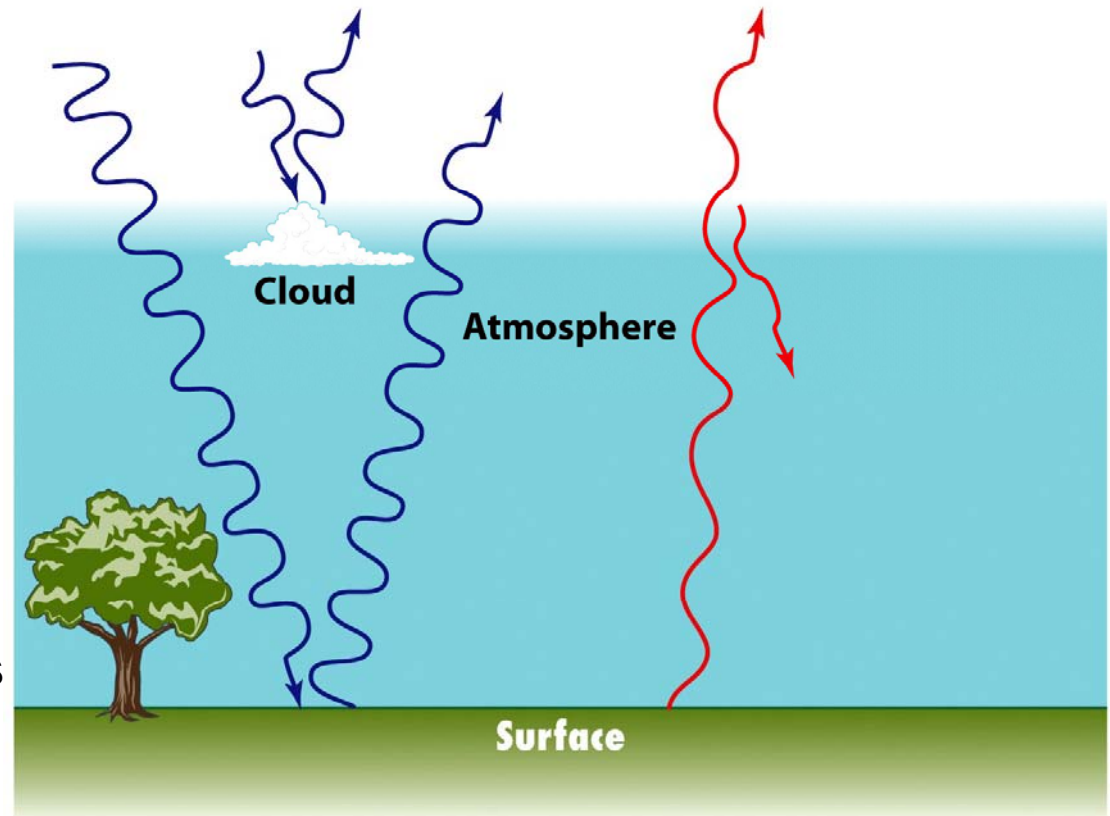
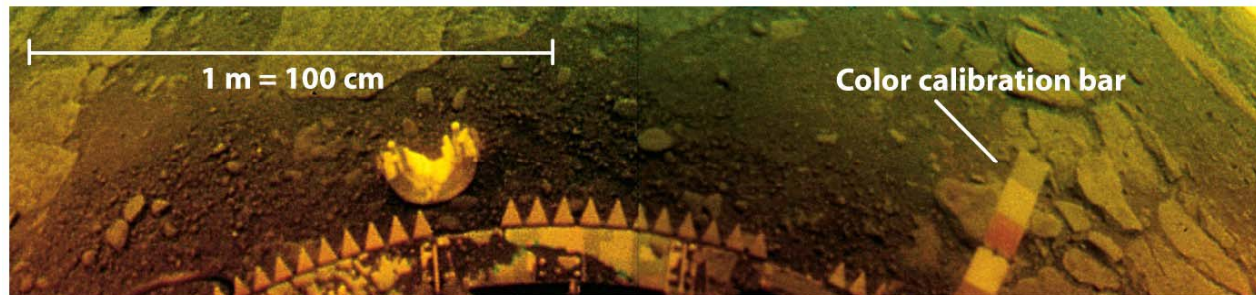
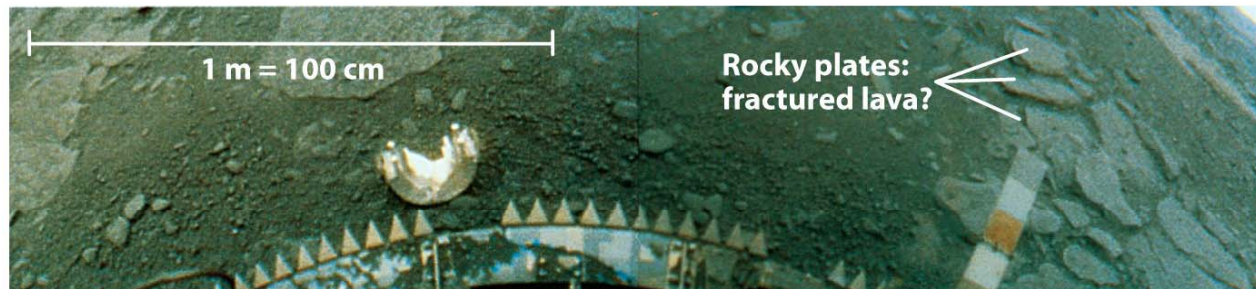


Figure 9-6  
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- **Venus and Earth's greenhouse effect compared**
  - **Earth's greenhouse boosts surface temperatures by about 33° C**
  - **Venus's greenhouse effects has boosted its surface temp. by 400° C**
  - **Result is that lead melts on the surface of Venus (and so do spacecraft)**
  
- **Venus also has highly reflective clouds**
- **If the clouds weren't there then the surface temperature would be 100s of degrees hotter**



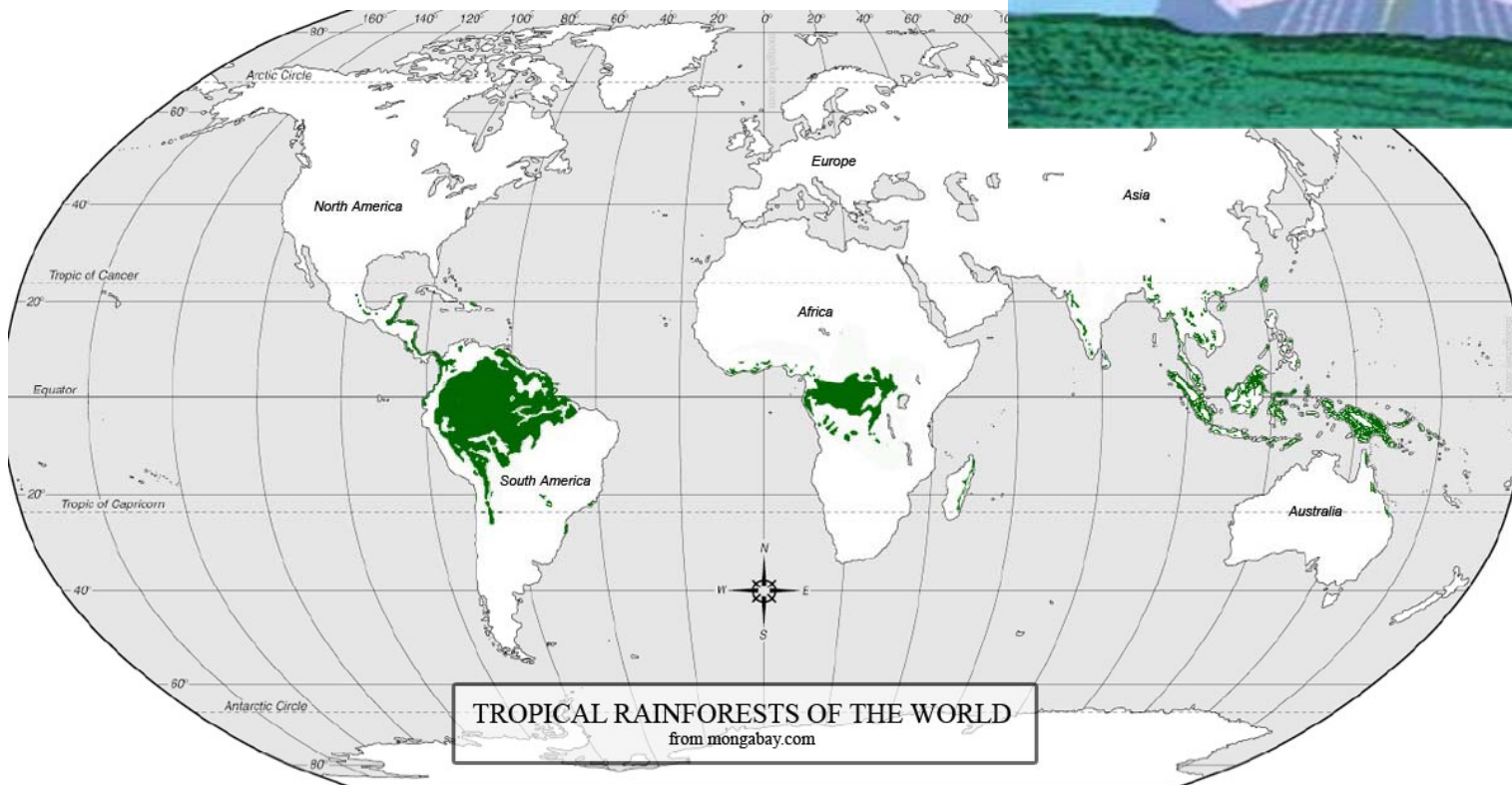
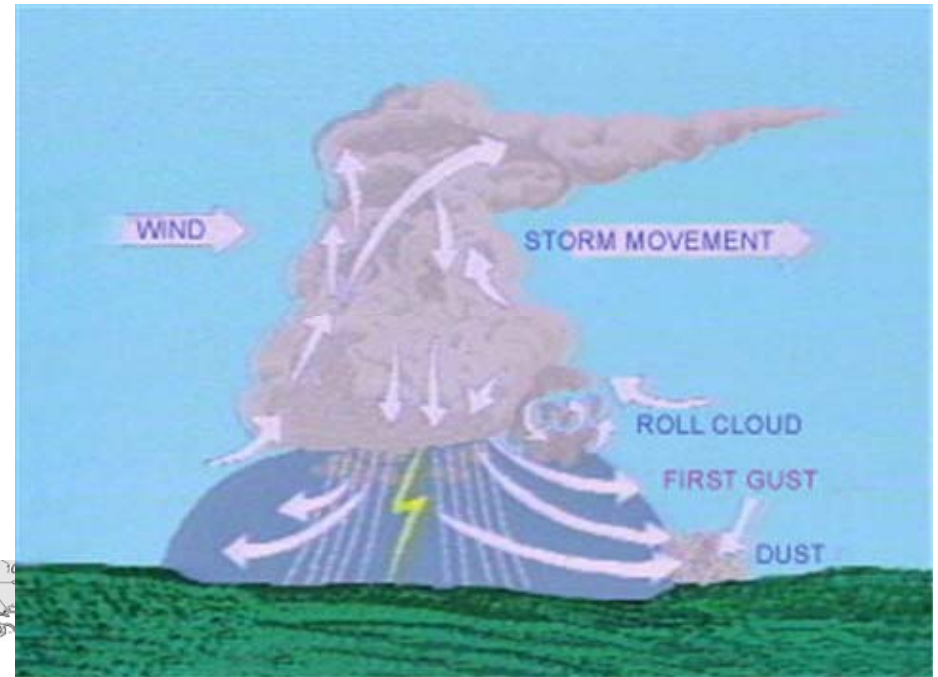
(a) Image from *Venera 13*



(b) Color-corrected image

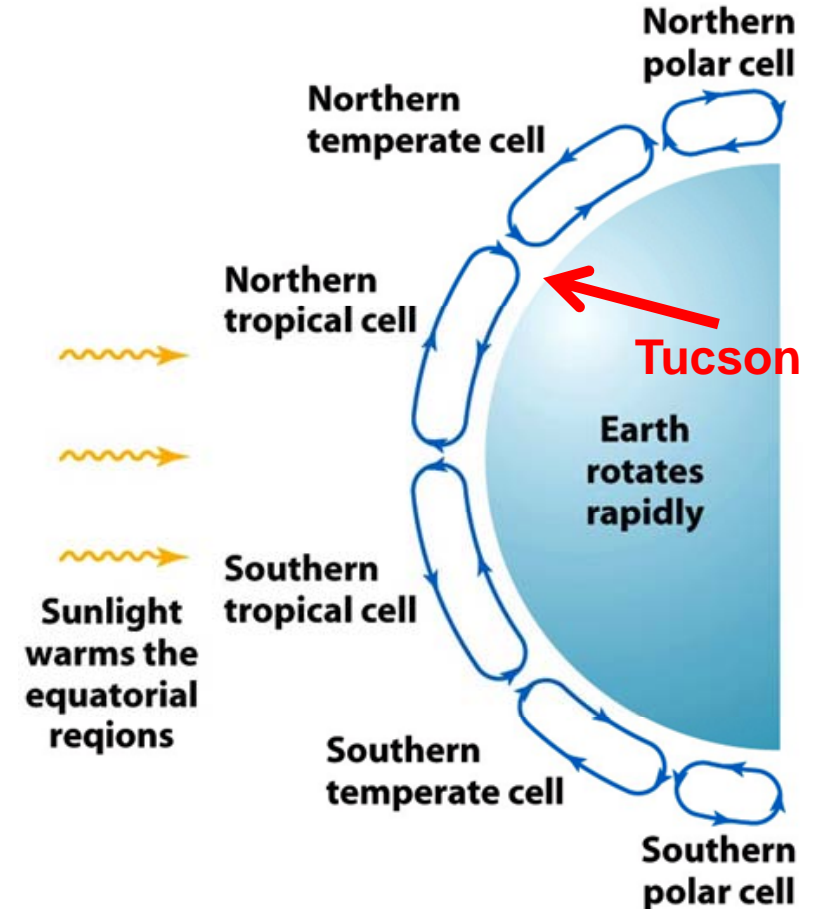
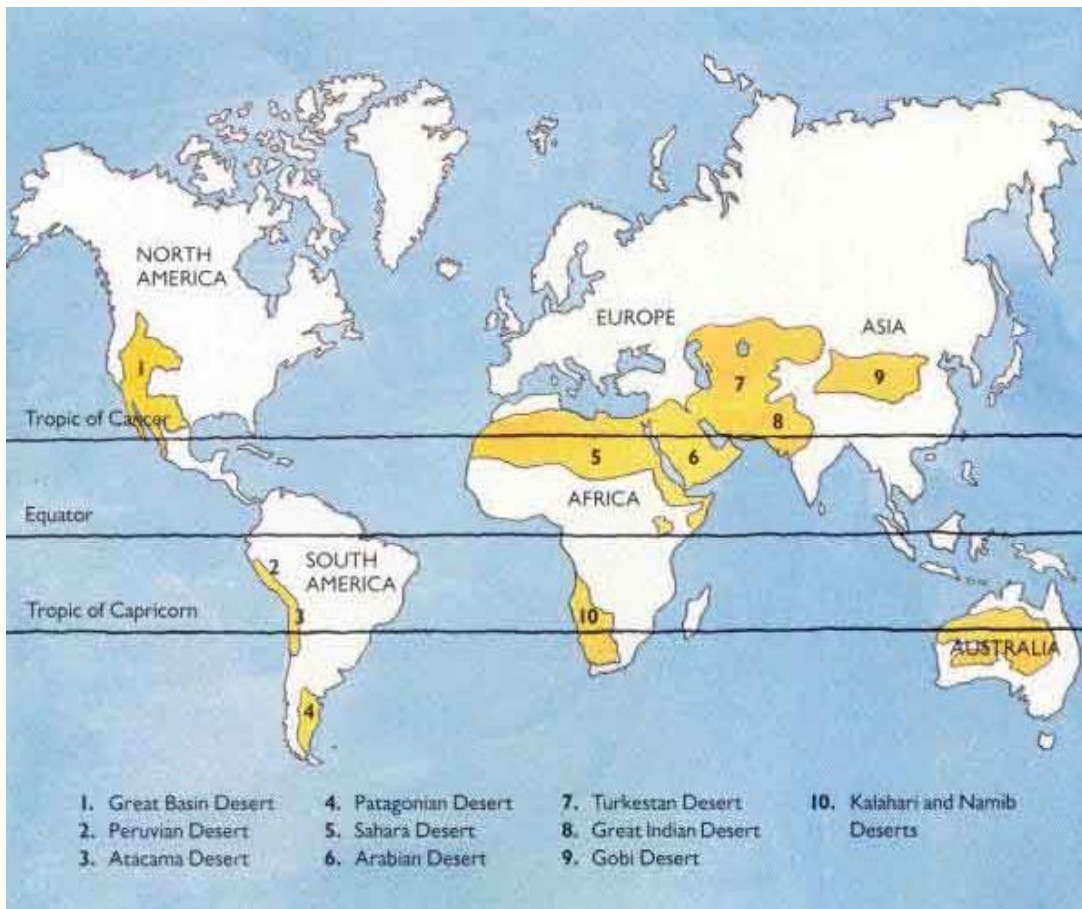
## Circulation

- Air gets heated at the equator
- Hot air rises and cools off
  - Clouds form and lots of rain results

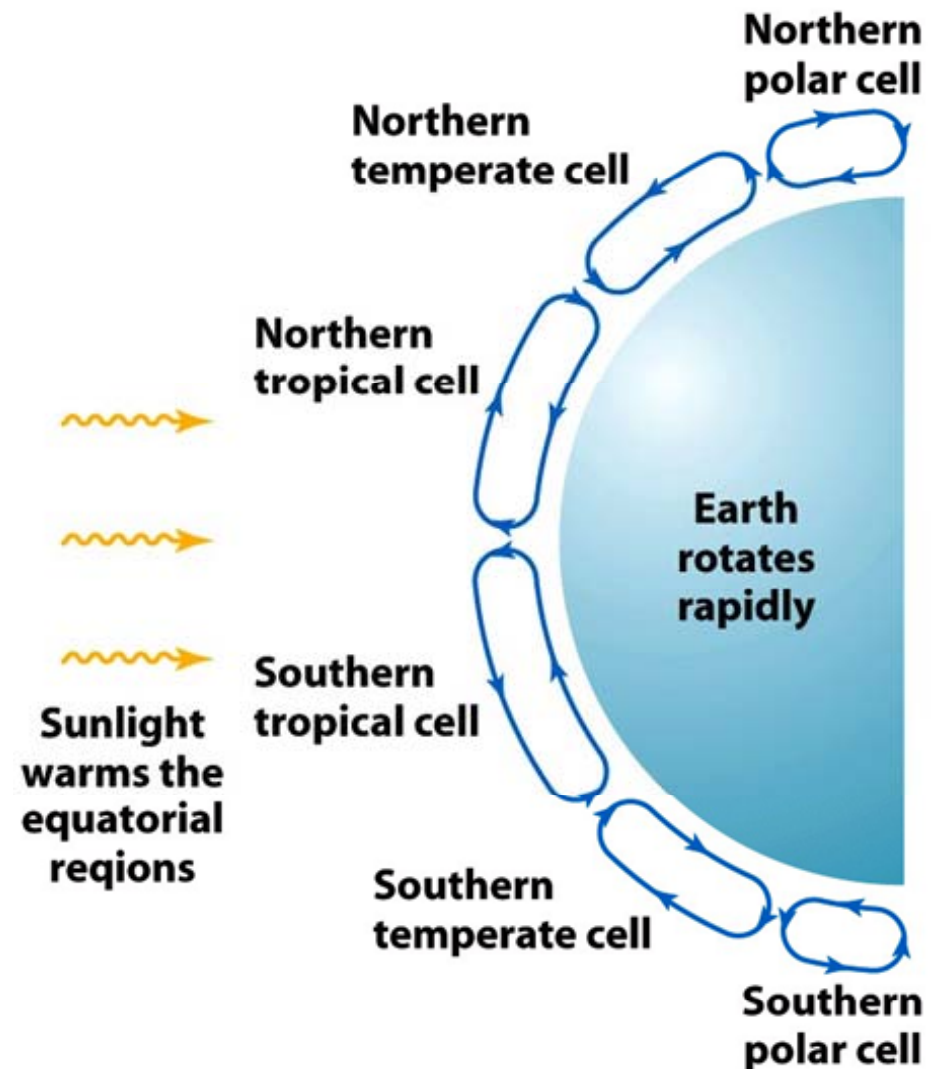


● **Hadley Cell circulation**

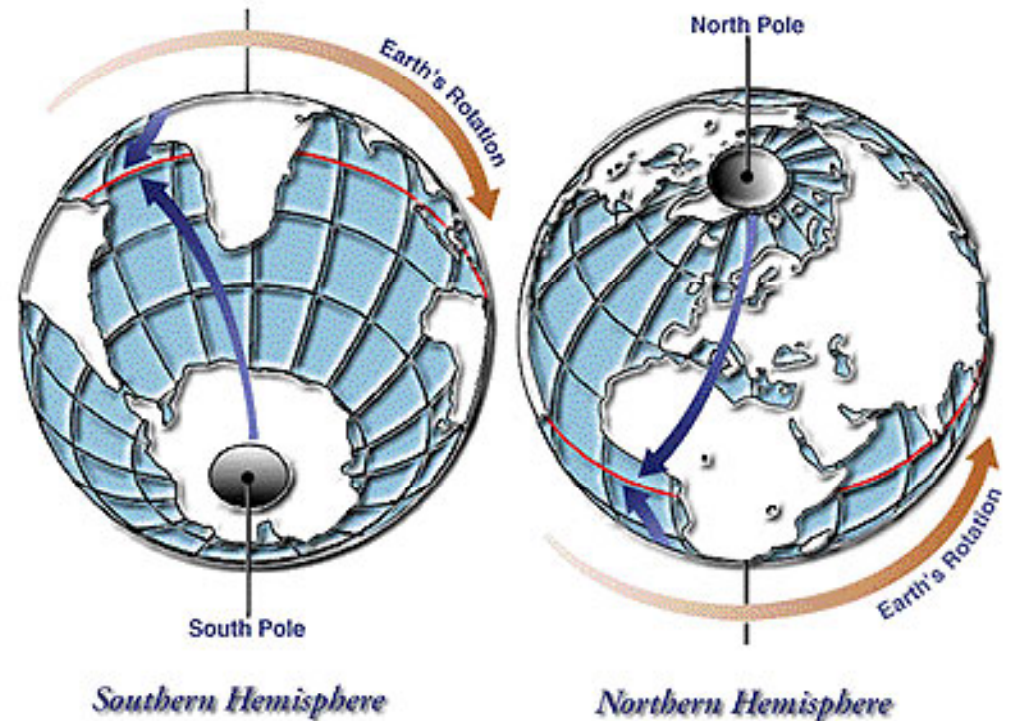
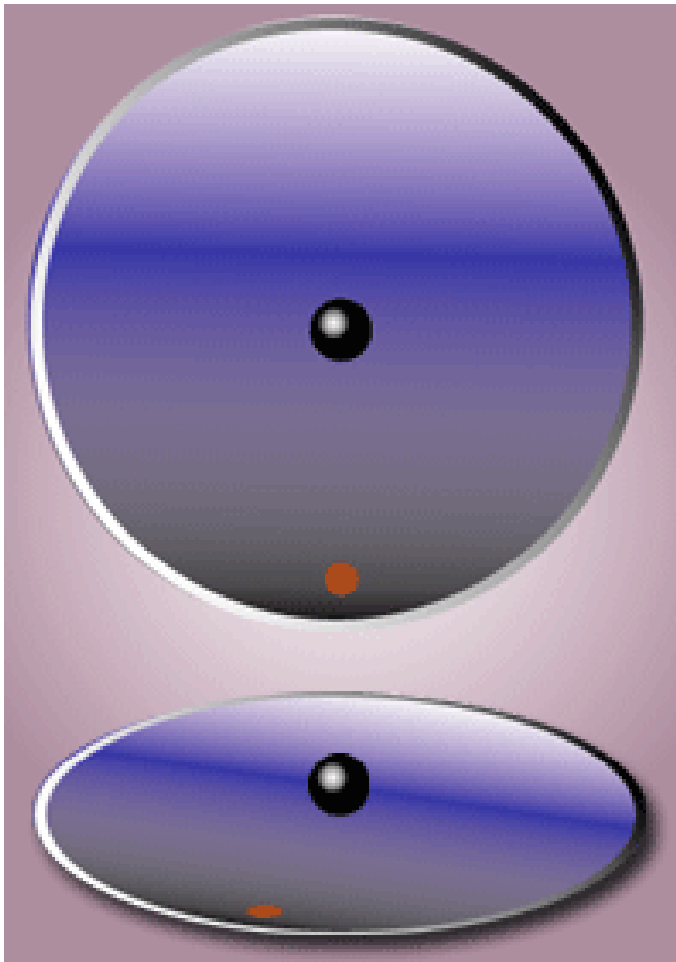
- Air (now dry) pushed aside from equator
- Moves to the tropics (~30° latitude)
- Descends to the surface and heats up – dry air creates deserts



- **Polar cell works the same way**
  - ...but in reverse
  - Cold air descends over the pole
  - Flows along the ground
  - Gets warmed up by being in contact with the surface
  - Rises around 60° latitude
  
- **Ferrell cell**
  - Sandwiched between the Hadley and polar cells
  - Driven by their motion

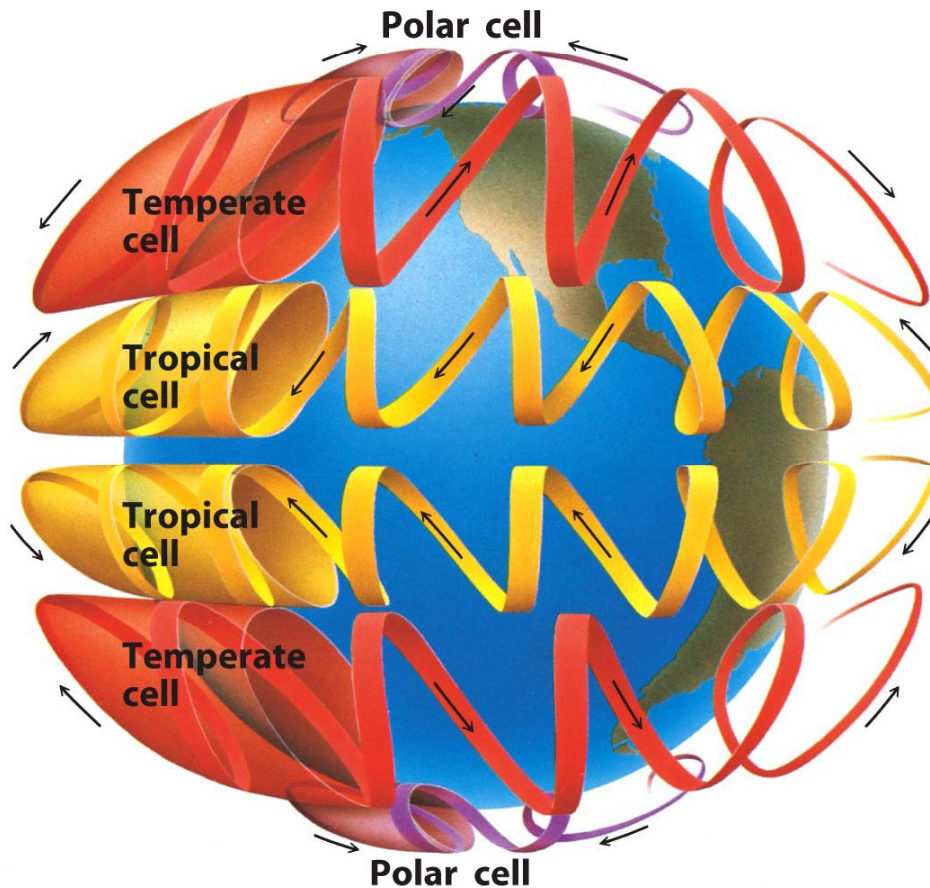


- What's the complication?
  - The Earth is spinning
  - Things that move don't travel in straight lines
  - Called the Coriolis effect



- In the Northern hemisphere
  - Things are deflectedd to the right
- In the Southern hemisphere
  - Things are deflectedd to the left

- This deflects moving air into zonal winds



**Doldrums – winds conflict**

**Trade winds – blow west at surface**

● Planetary comparison?

- Venus has a huge Hadley cell
- Atmosphere is so thick that temperatures don't vary much between the pole and equator (or between day and night)
- Rotation rate is very low so not much Coriolis deflection

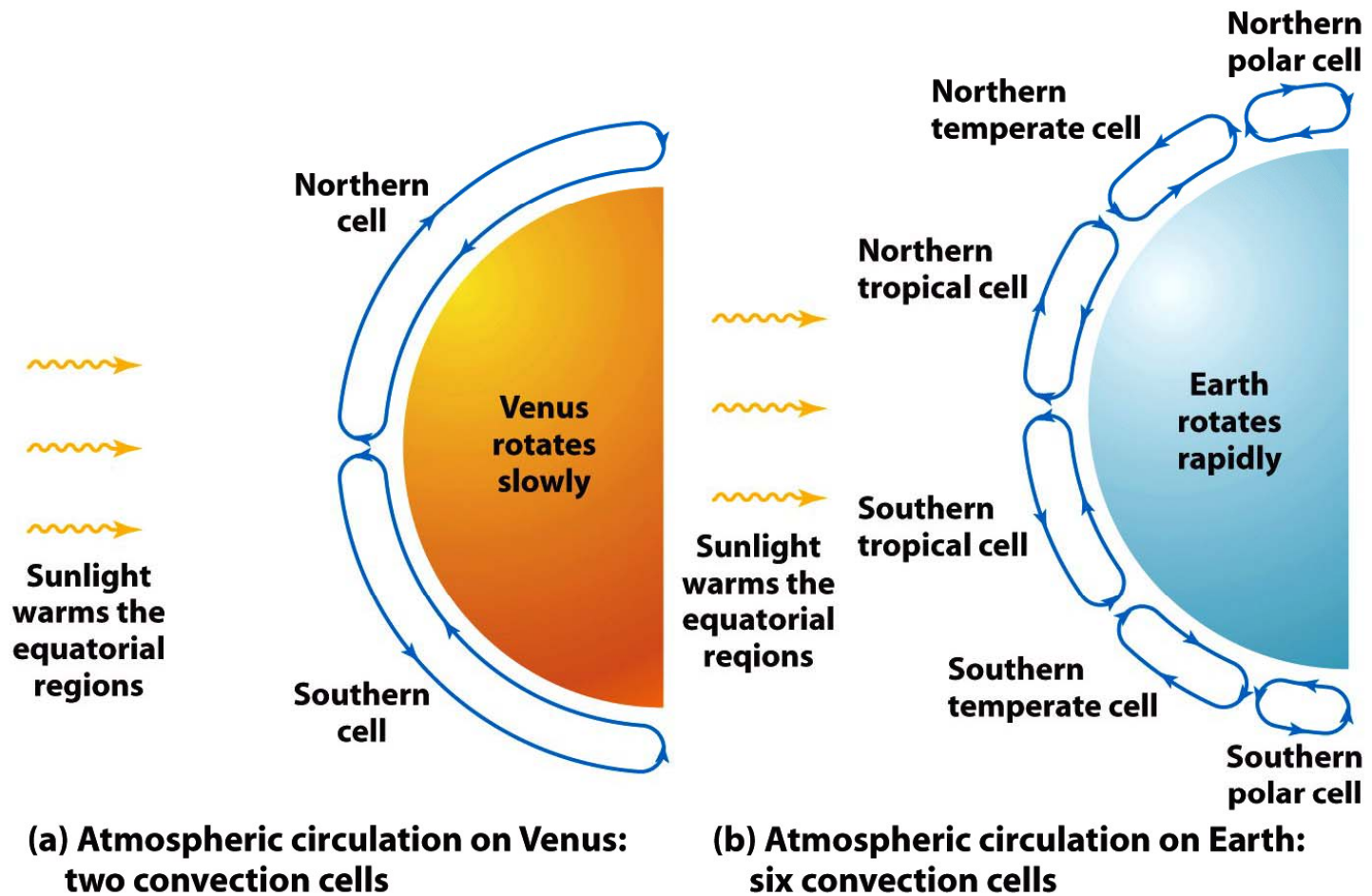


Figure 11-28  
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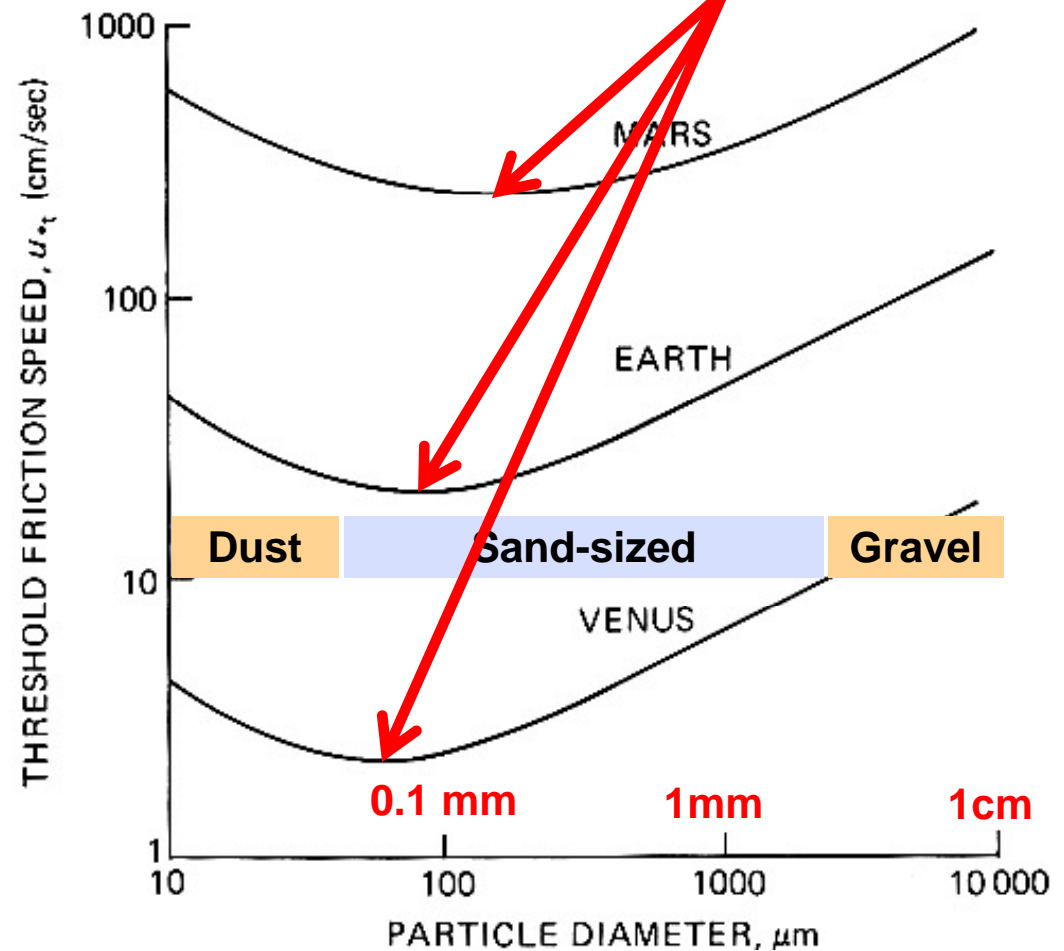


## Surface features from the atmosphere

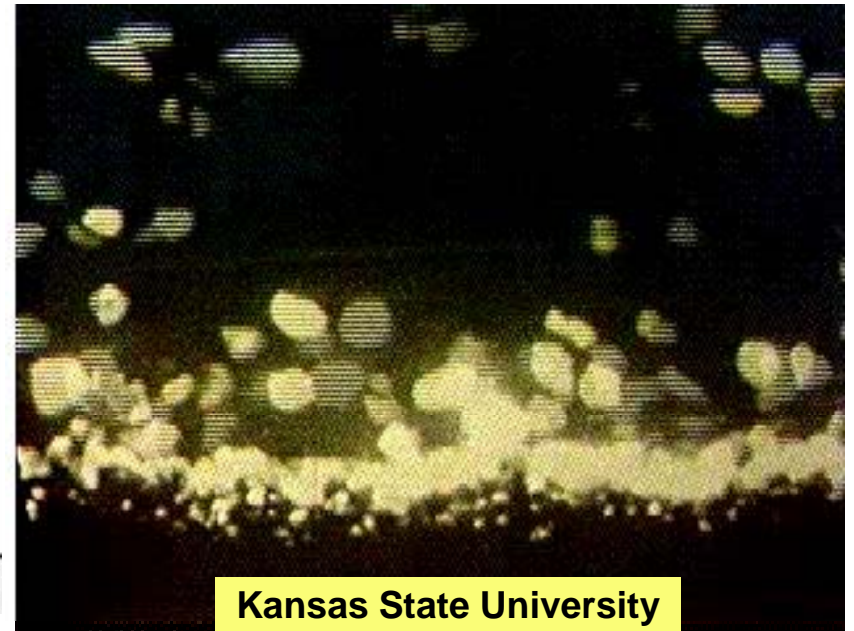
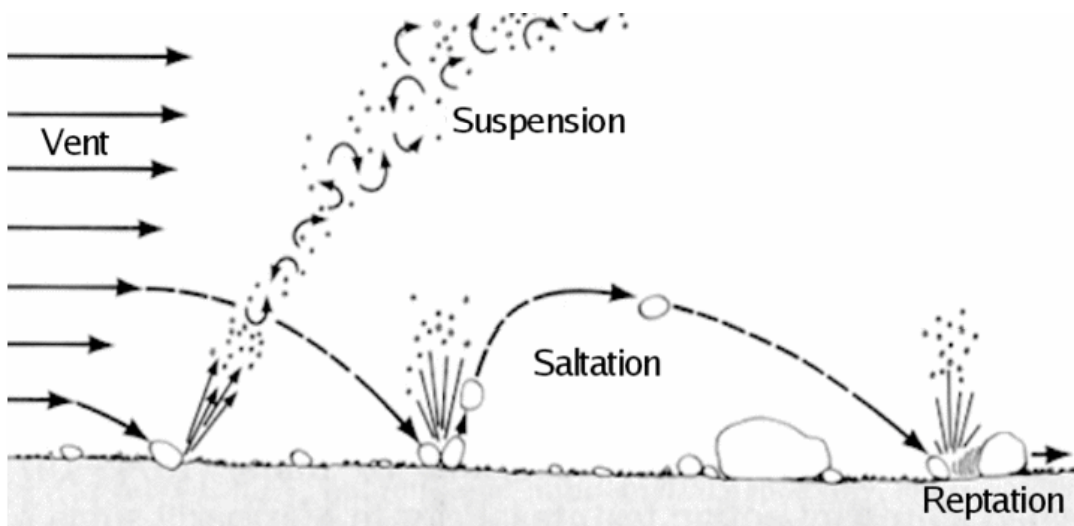
- Particles classified by Udden-Wentworth scale

$\phi$ scale	Size range (metric)	Aggregate name (Wentworth Class)
< -8	> 256 mm	Boulder
-6 to -8	64–256 mm	Cobble
-5 to -6	32–64 mm	Very coarse gravel
-4 to -5	16–32 mm	Coarse gravel
-3 to -4	8–16 mm	Medium gravel
-2 to -3	4–8 mm	Fine gravel
-1 to -2	2–4 mm	Very fine gravel
0 to -1	1–2 mm	Very coarse sand
1 to 0	½–1 mm	Coarse sand
2 to 1	¼–½ mm	Medium sand
3 to 2	125–250 $\mu\text{m}$	Fine sand
4 to 3	62.5–125 $\mu\text{m}$	Very fine sand
8 to 4	3.90625–62.5 $\mu\text{m}$	Silt
> 8	< 3.90625 $\mu\text{m}$	Clay
>10	< 1 $\mu\text{m}$	Colloid

**Easiest particles to move are sand-sized**

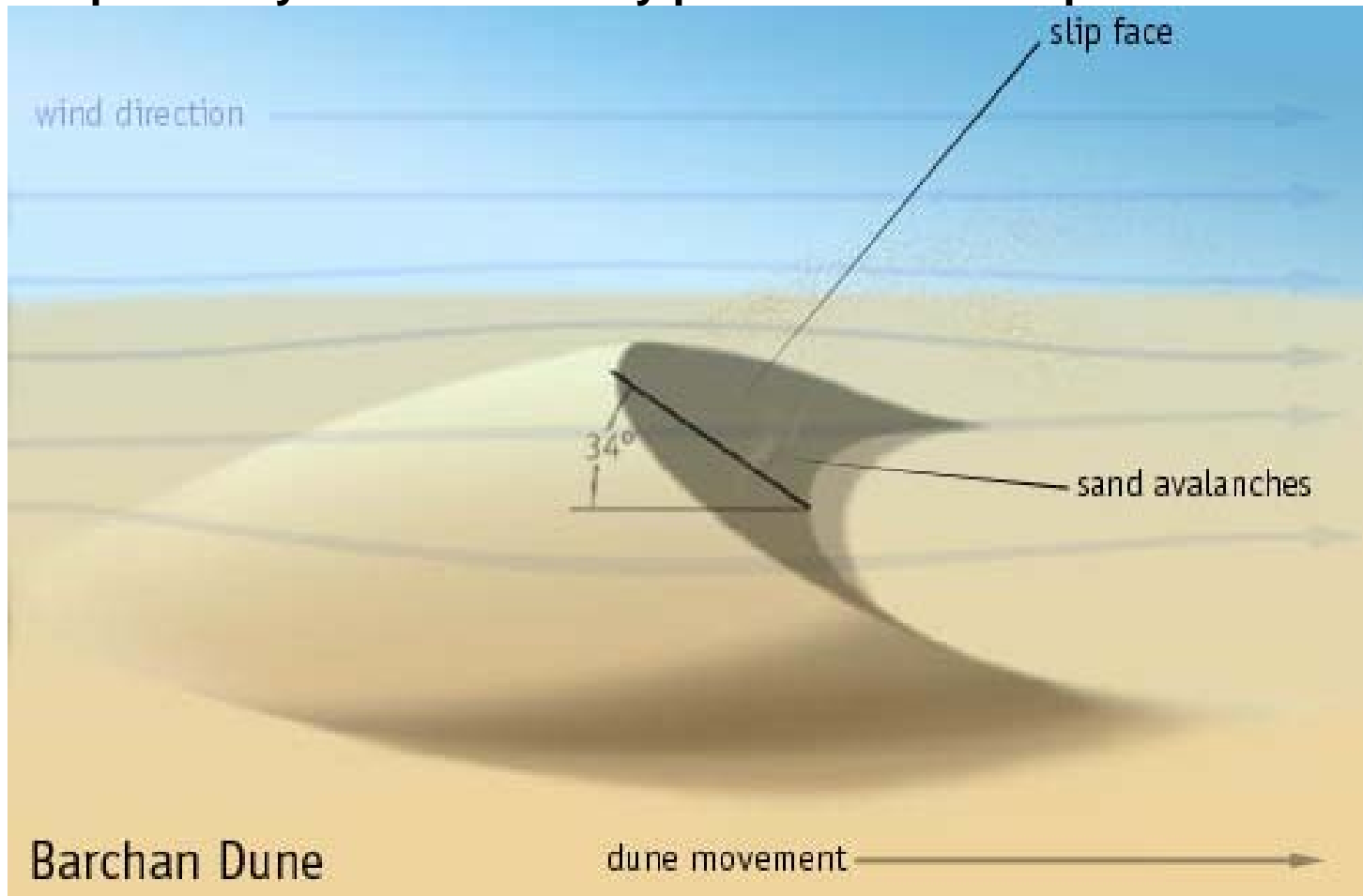


- Grains travel by saltation
  - Impacting grains can dislodge new particles (reptation)
  - Impacting grains can push larger particles (creep)



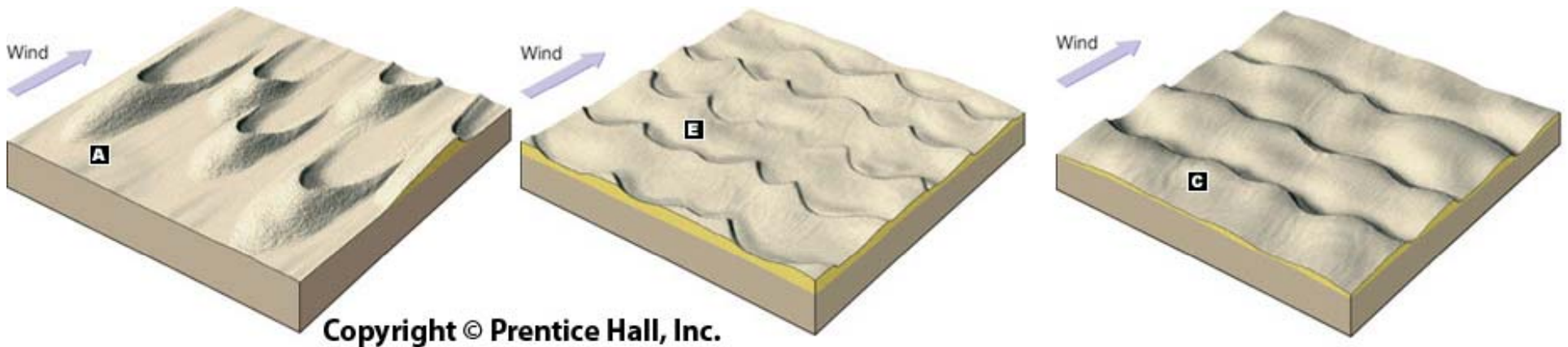
Kansas State University

- Saltating particles make many hops up the windward slope
- Wind disconnects from the surface at edge of dune
- Sand falls down the ‘slip-face’
  - Slip face slope is the angle of repose
  - Angle of repose very similar for many particle-sizes and planets



- Unidirectional winds

- Dune shape depends on sand supply

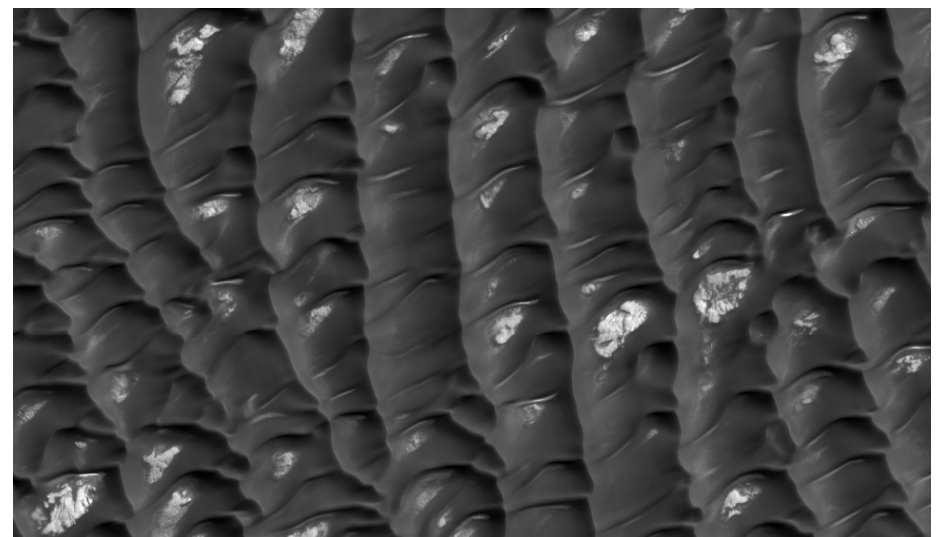
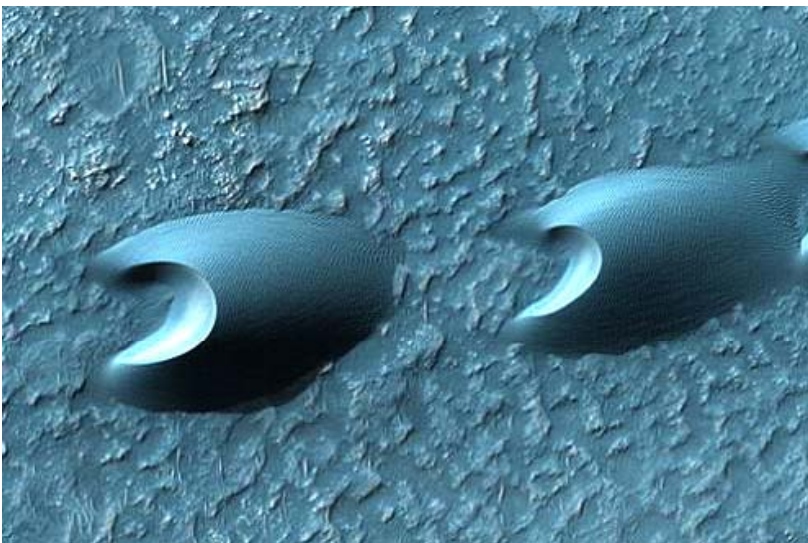


Barchan

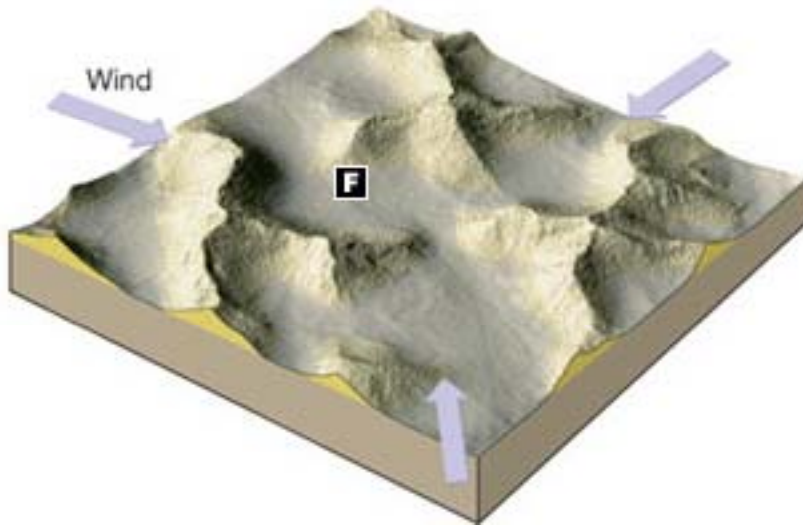
Barchanoid ridges

Transverse

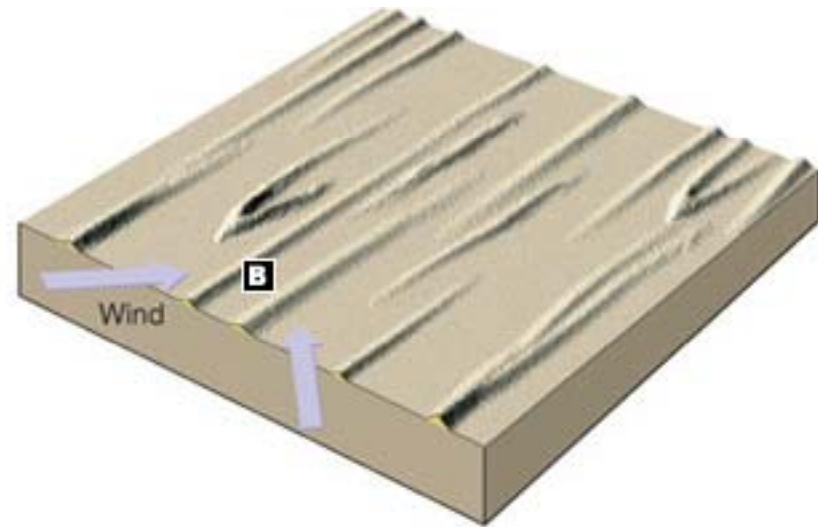
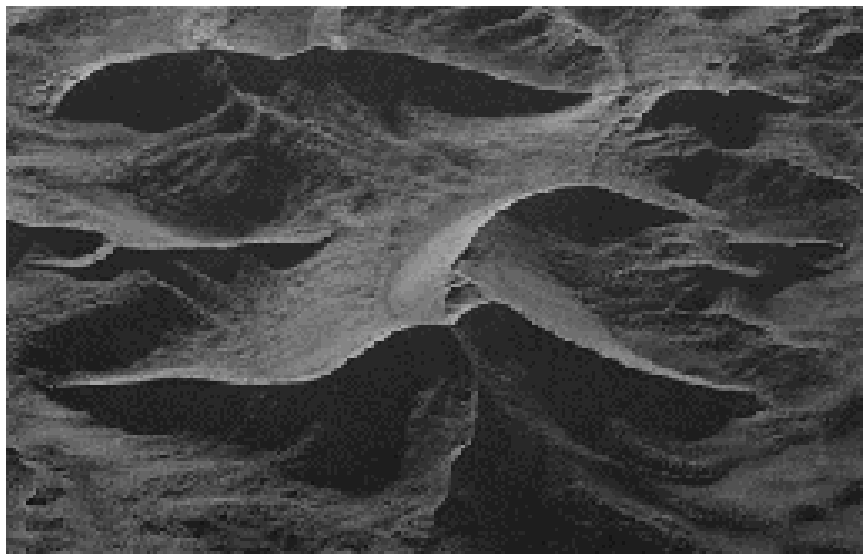
Increasing sand supply 



- When wind directions vary
  - Linear dunes – a little variation
  - Star dunes – a lot of variation



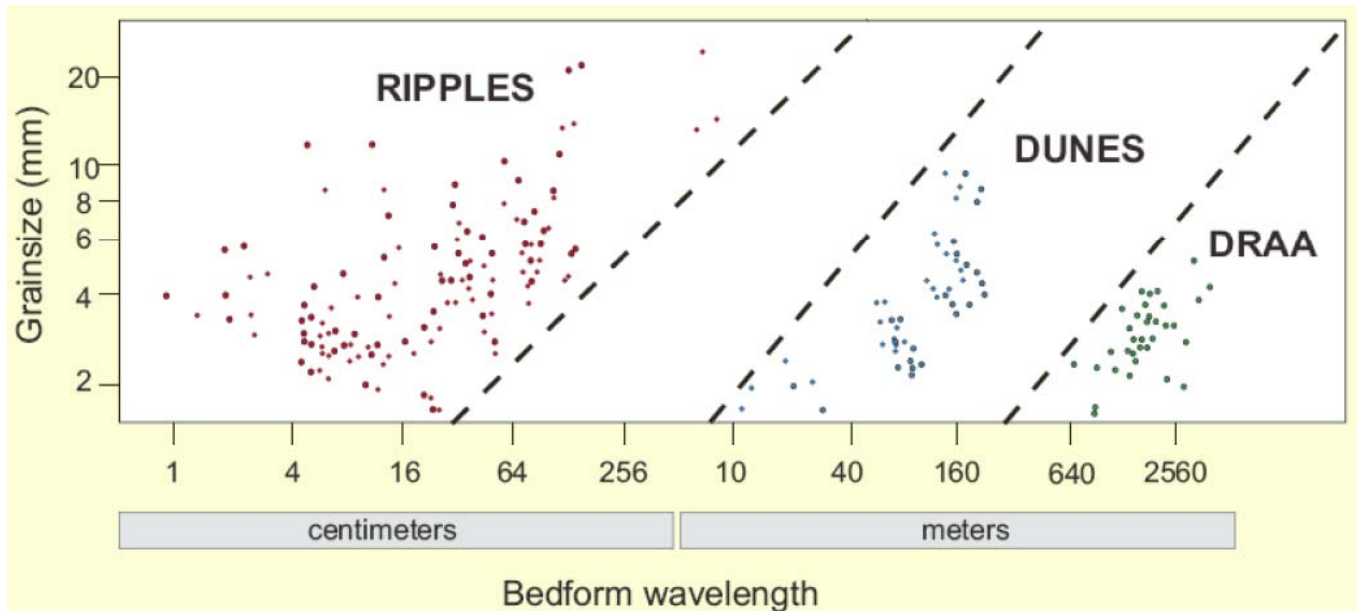
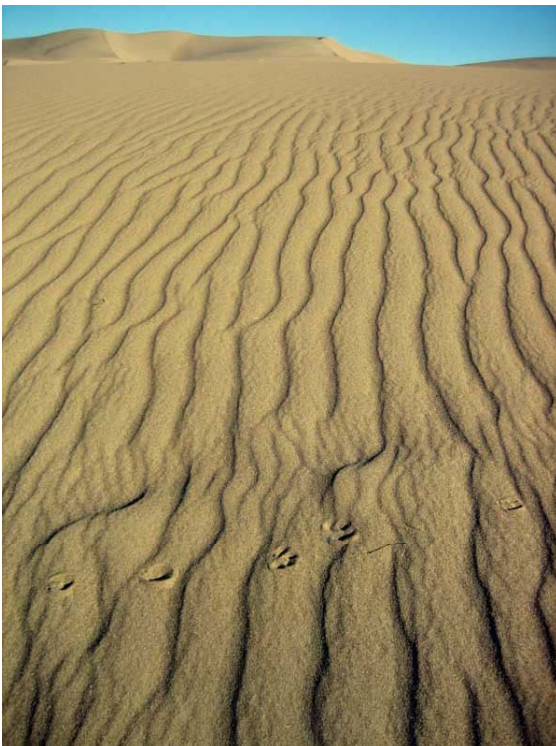
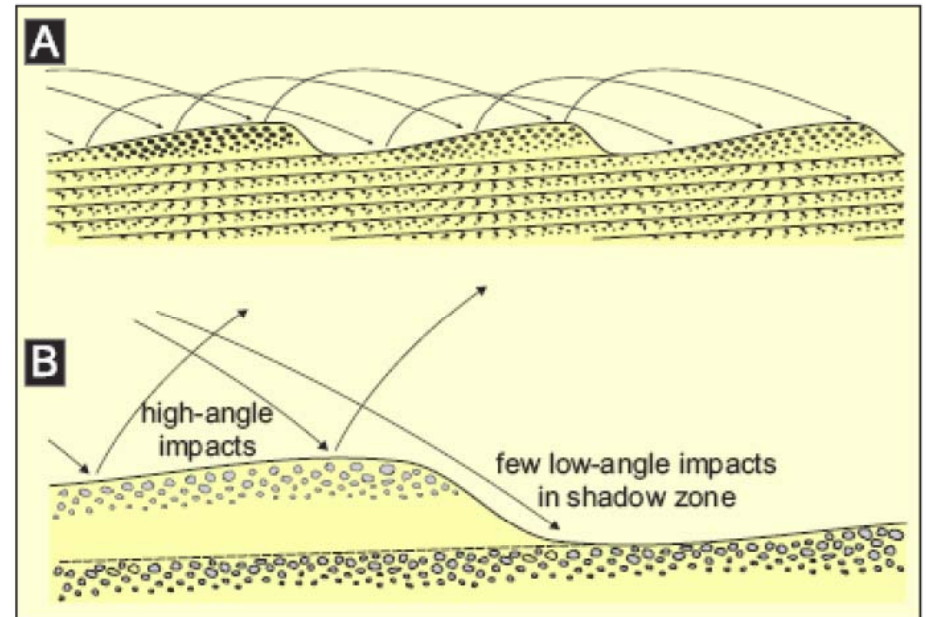
Star



Linear/Longitudinal



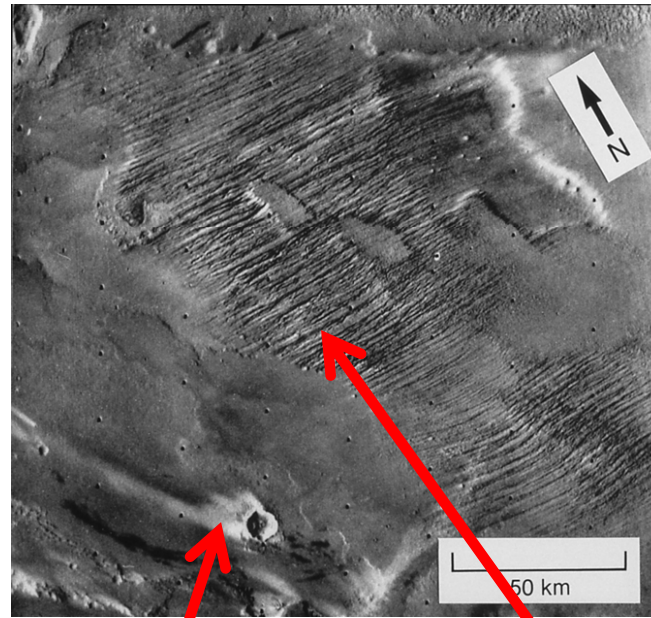
- Saltating particles can also form ripples
  - Much more dynamic than dunes
  - Asymmetric, 8-10 vs 20-30 degree slopes
  - Typical wavelengths 7-14cm, heights 0.5-1cm



**Aeolian Erosion**

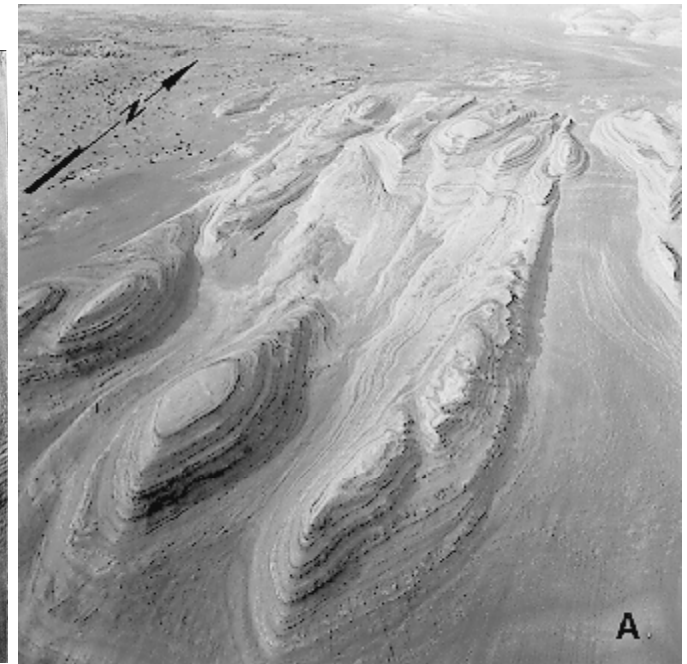
- **Yardangs**

- Wind blown particles abrades surface
- Erosion leaves elongated mounds as remnants
- Requires strong, virtually unidirectional, wind.
- Eroded material must be consolidated



Recent wind action

Older wind action

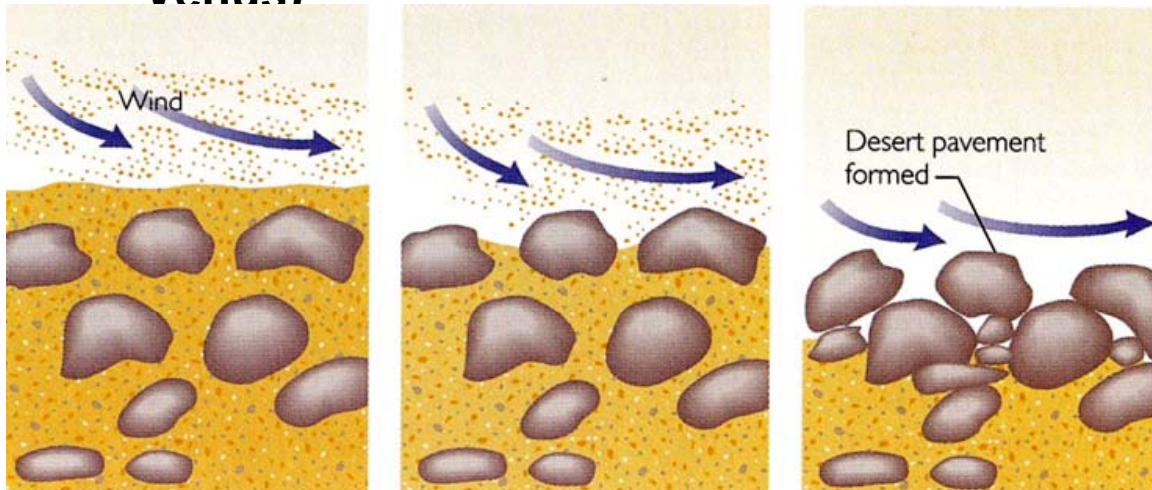


- **Ventifacts**

- Elongated erosional marks on rocks
- Usually works on originally circular vesicles
- Used as paleo-wind direction indicators e.g. pathfinder landing site

- **Desert Pavement**

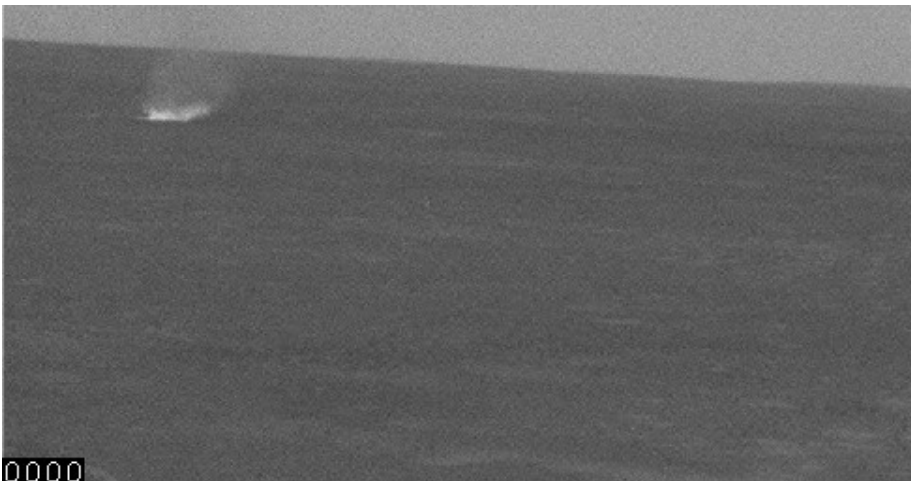
- Occurs frequently on Earth and Mars (and probably Venus)





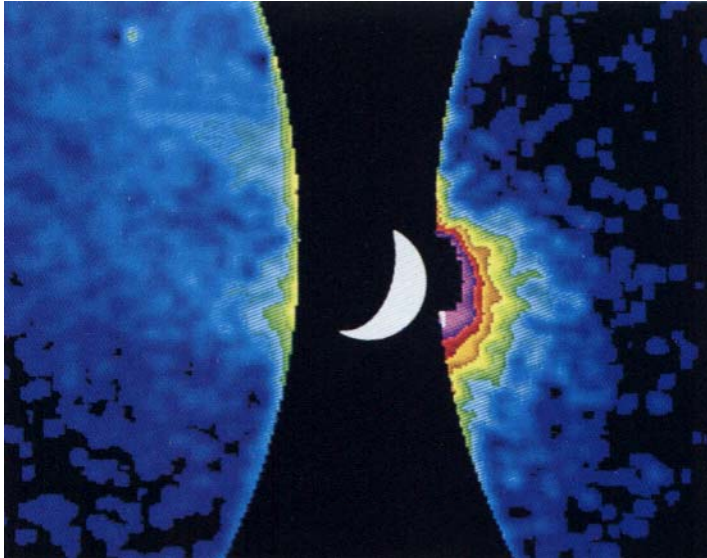
## Dust Devils

- **Plume of hot air rises**
- **Penetrates cooler atmospheric layer**
  - gets stretched out vertically
  - Conservation of angular momentum starts column spinning
  - More hot air gets sucked in at the base to keep it self-sustaining



## 'Atmospheres' of the Moon and Mercury

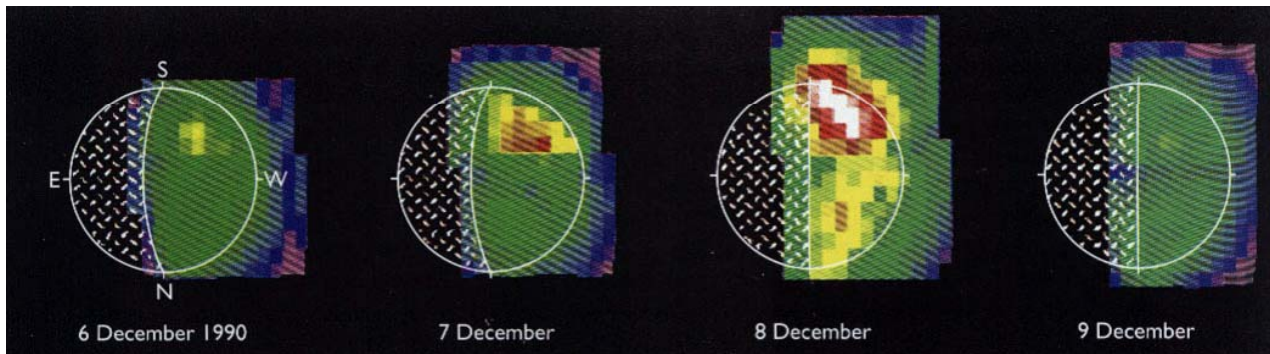
- The Moon and Mercury both possess tenuous atmospheres



### Two Tenuous Atmospheres

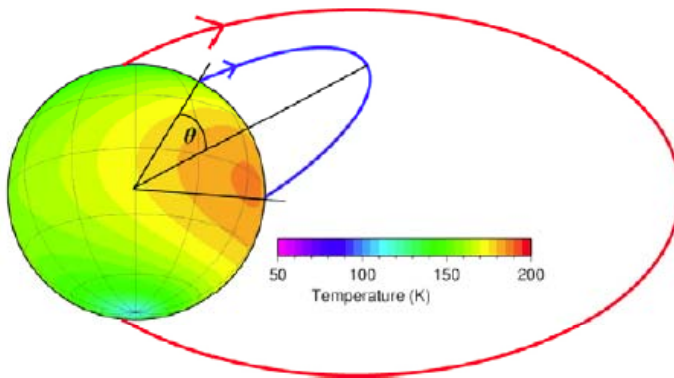
Species	Mercury (atoms/cm <sup>3</sup> )	Moon (atoms/cm <sup>3</sup> )
Hydrogen (H)	200	< 17
Helium (He)	6,000	2,000–4,000
Oxygen (O)	< 40,000	< 500
Sodium (Na)	20,000	70
Potassium (K)	500	16
Argon (Ar)	< 3 × 10 <sup>7</sup>	4 × 10 <sup>4</sup>

Calcium now also seen at Mercury



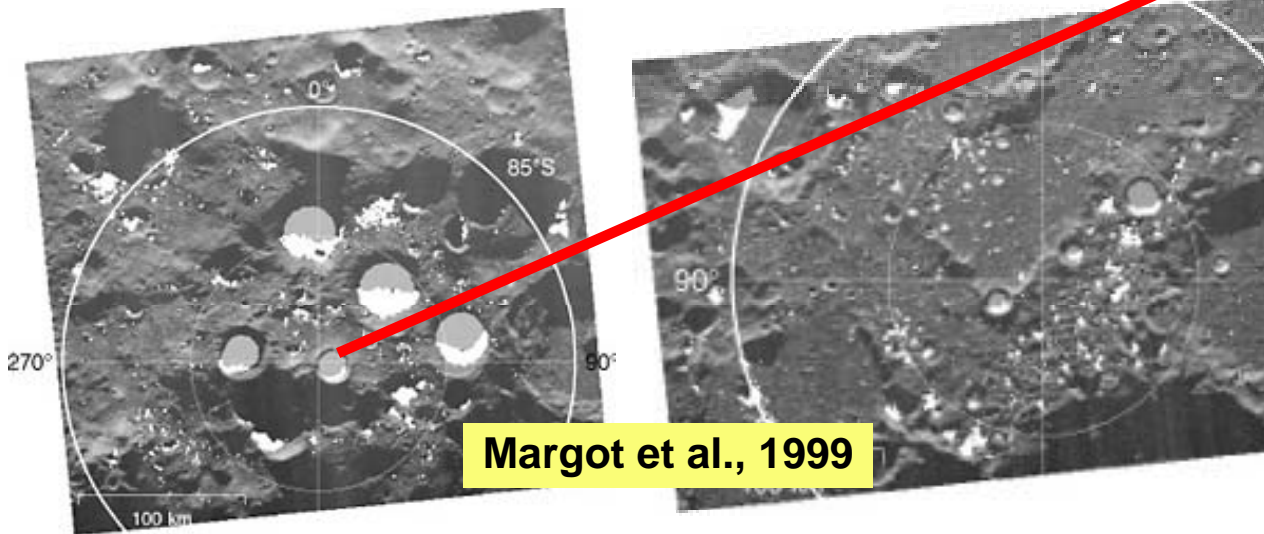
- Sodium emission at the Moon and Mercury shows temporal changes
  - Stirring of regolith by small impacts

- **Airless bodies do have ‘atmospheres’**
  - **Surface bounded exospheres**
  - **Atoms collide more often with the surface than with each other**  
 mean free path  $\gg$  atmospheric scale height  
 (really means that mean free path  $\gg$  trajectory of a molecule)

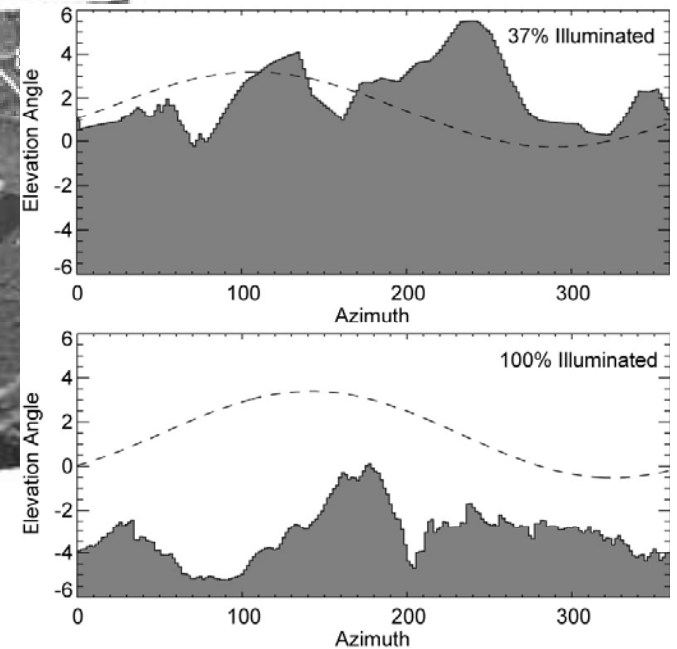


- **Particles hop around until they find cold spots (e.g. night-side or shadowed area)**
  - ◆ Ejection rate is slow & range is small
- **When the sun comes up they start hopping around again**

- Do permanently shadowed regions exist?
  - Yes, Moon and Mercury have low obliquity
  - Solar elevations in the polar regions are always low
  - Even modest craters can have permanent shadow on their floors

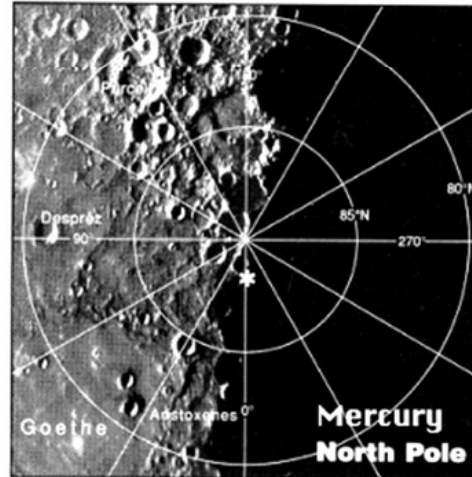
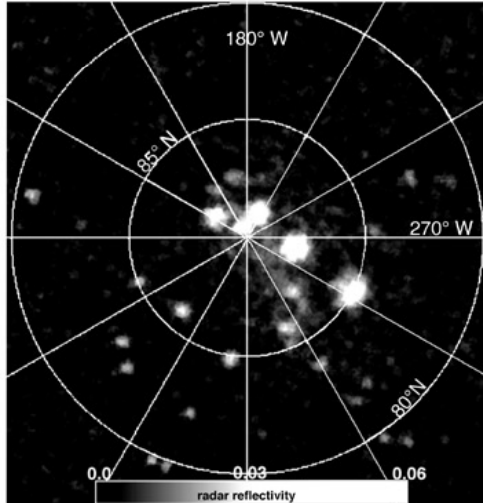


Margot et al., 1999

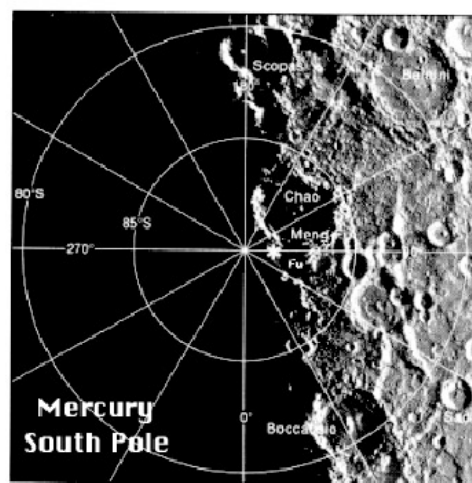
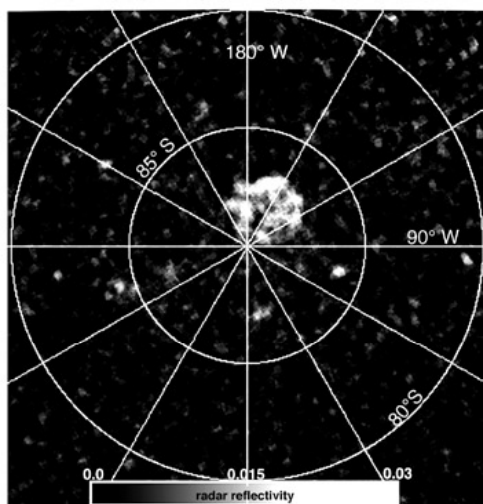


- Evidence for ice in polar craters of the Moon and Mercury
  - Radar reflections show craters are filled with unusual material
  - Probably water ice

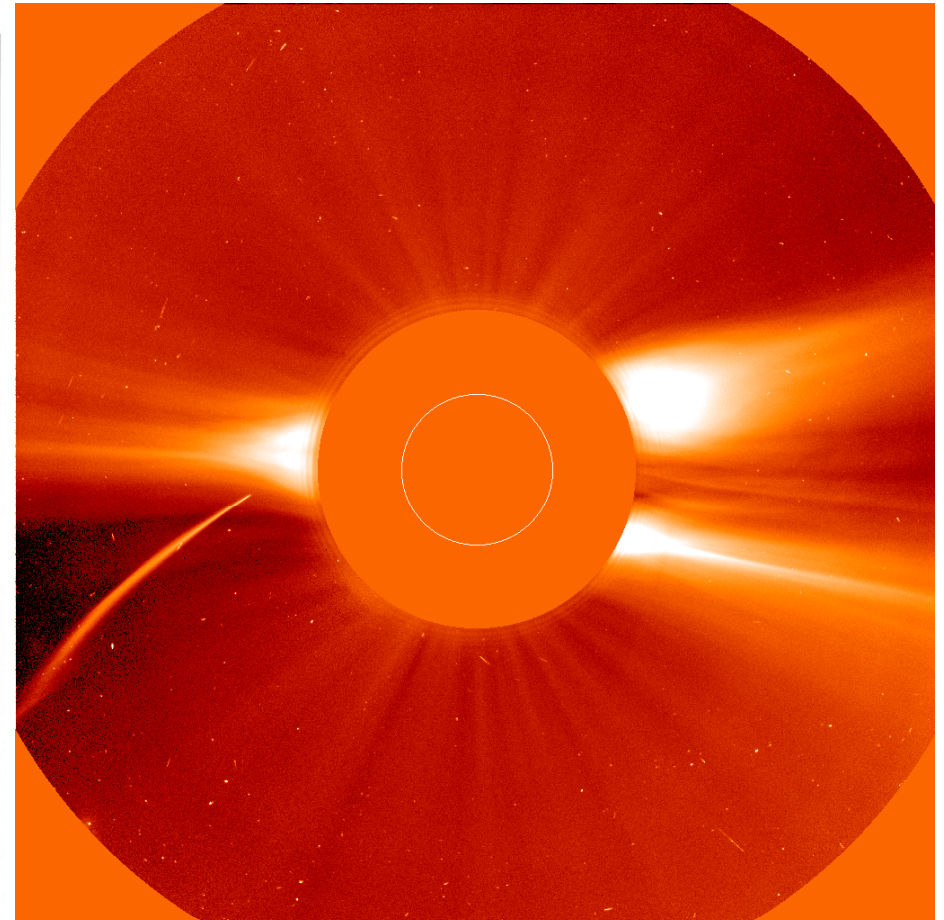
MERCURY NORTH POLAR ARECIBO RADAR IMAGE



MERCURY SOUTH POLAR ARECIBO RADAR IMAGE

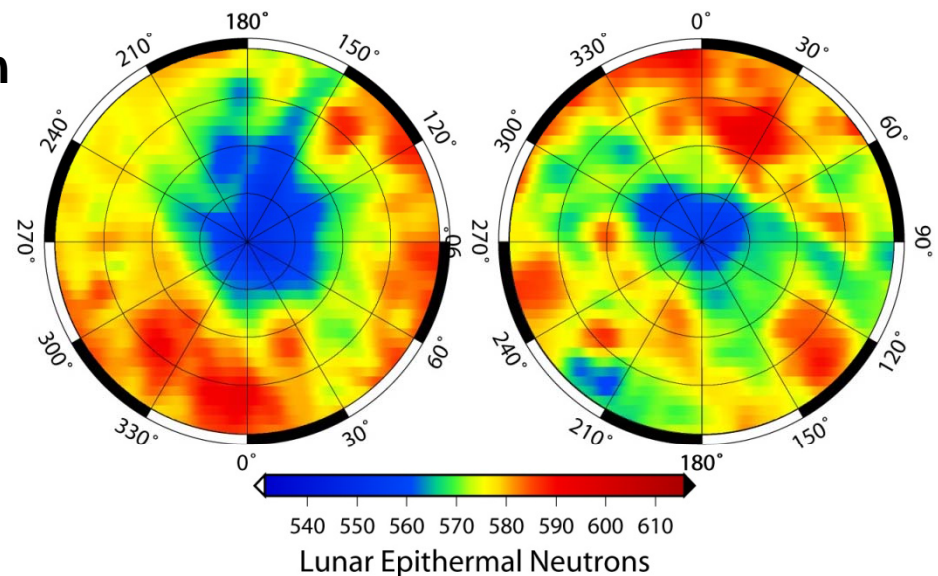
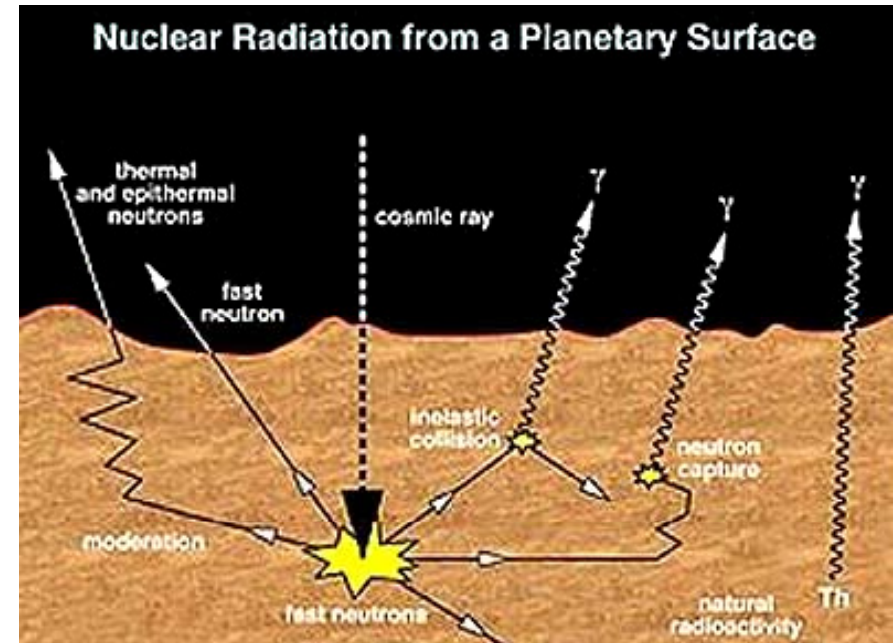


(Courtesy J. K. Harmon and M. A. Slade)



- Source of the water?
  - Kreutz group – sungrazing comets

- **Neutron spectrometer data**
  - **Cosmic rays hit the surface**
  - **Produce neutrons**
  - **Neutrons are easily scattered by hydrogen**
  
- **Areas with low neutron counts have a lot of hydrogen**
- **These areas correspond to the polar areas on the Moon**
  
- **Hydrogen usually means water (ice in this case)**
  
- **Clementine bi-static radar**
  - **This evidence is pretty weak**
  
- **Lunar poles appear most empty of ice**
  - **No recent impacts?**





## In this lecture...

- **Introduction to atmospheres**
  - Pressure, Temperature, Scale height
- **Radiation and atmospheres**
  - Clouds & Greenhouse effect
- **Circulation**
- **Surface features from the atmosphere**
  - Sand dunes and how they work
  - Eolian erosion
- **'Atmospheres' of Moon and Mercury**
  - Ice in polar craters

## Next: Venus

- **Reading**
  - Chapter 11-6 & 11-7 to revise this lecture
  - Chapter 11-Venus sections for next lecture